

# **CARBON-BEARING VOLATILES:**

**SURFACE ABUNDANCE ESTIMATES FROM  
EXOSPHERIC CONTENT CONSIDERATIONS.**

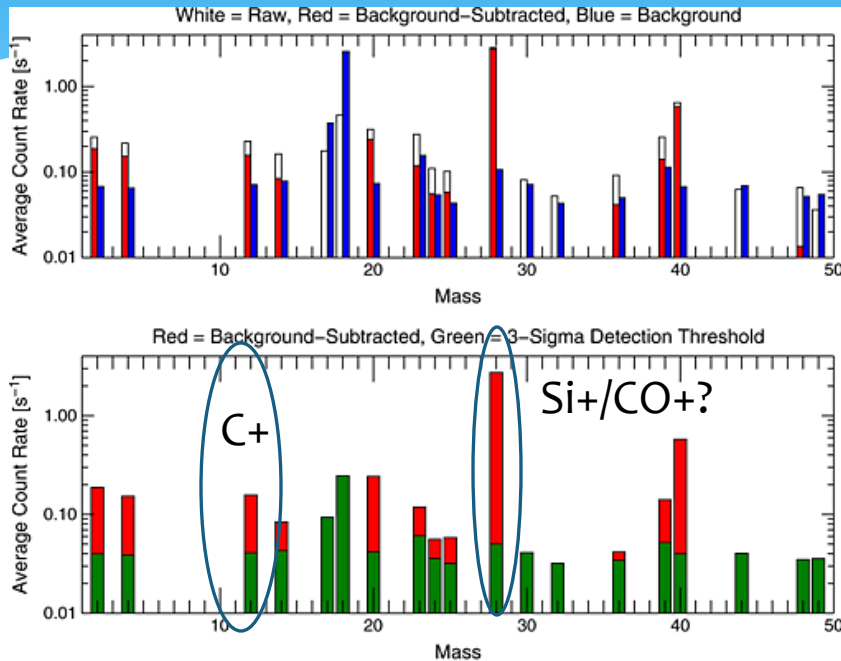
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# An exosphere of CO and CO<sub>2</sub>?

## Detections of lunar exospheric ions by the LADEE neutral mass spectrometer



Halekas et al. GRL 2015

If [CO<sup>+</sup>], then [CO] ~ 10<sup>4</sup> cc

- Ion Measurements have long indicated predominance of mass 28 ions (Al<sup>+</sup>, Si<sup>+</sup>, or CO<sup>+</sup>?)


### Motivating Questions:

- ☐ What are the expected levels of carbon-bearing volatiles in the exosphere of the Moon?
- ☐ And, based on flux balance, what would be the distribution of adsorbed species on the surface of the grains?

# Expected carbon influx as the seed population for these exospheres

## 1) Solar wind:

- A long-term average of the carbon to proton abundance in the solar wind is  $\sim 2 - 3 \times 10^{-4}$  [von Steiger et al., 2000].
- Assuming the average solar wind flux to be  $2 \times 10^8$  ions  $\text{cm}^{-2} \text{s}^{-1}$ , the carbon influx to the Moon is  $4 - 6 \times 10^4$  ions  $\text{cm}^{-2} \text{s}^{-1}$



Methane gas,  
Trapped {carbides,  
CO, CO<sub>2</sub>}

## 2) Micrometeoroids:

- The carbon fraction in CI chondrites is 8.5 wt% [Lodders and Fegley, 1998].
- Adopting a lower limit of the rate of micrometeoritic flux onto the Moon of  $5.12 \times 10^{-17}$  gm  $\text{cm}^{-2} \text{s}^{-1}$  [Cintala, 1992] and an upper limit of  $4.76 \times 10^{-16}$  gm  $\text{cm}^{-2} \text{s}^{-1}$  (Furi et al., 2012) the carbon influx, F, is:  $8.8 \times 10^4 < F < 8.2 \times 10^5$  C atoms  $\text{cm}^{-2} \text{s}^{-1}$



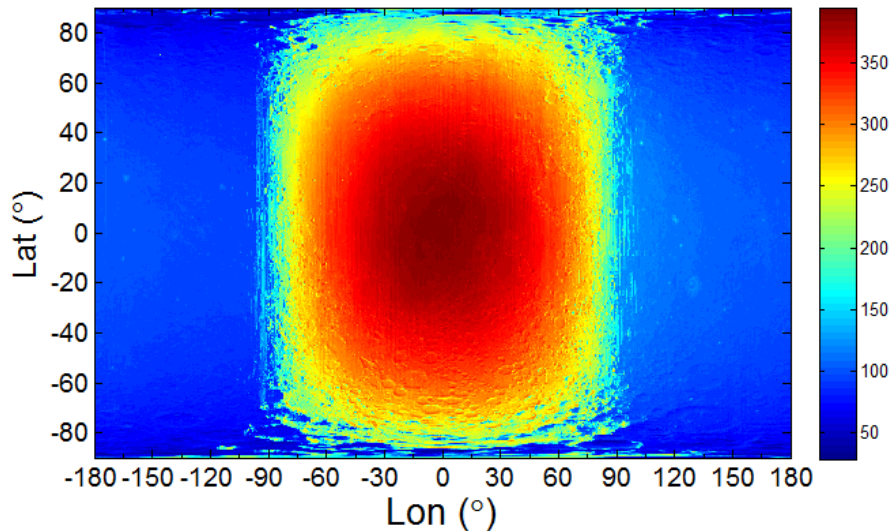
CO, CO<sub>2</sub> gas

# Mobility of carbon volatiles

Adsorbate	Adsorbent/Surface	Eact (eV)	Source
CO	TiO <sub>2</sub> (100)	0.42	Linsebigler et al 1995
CO	TiO <sub>2</sub>	0.45-0.5	Raupp and Dumesic 1985
CO	ZnO (powder)	0.36	Wang et al. 2007
CO coadsorbed with CO <sub>2</sub>	ZnO (powder)	0.64	Wang et al. 2007
CO	Fe (100)	0.62,0.88,1.08	Benziger and Madix, 1980
CO <sub>2</sub>	MgO (100)	0.40	Meixner et al. 1992
CO <sub>2</sub>	ZnO (powder)	0.46-0.67	Xia et al. 2008
CO <sub>2</sub>	TiO <sub>2</sub>	0.46	

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- Major knowledge gap!
  - Co-adsorbates influence the binding
    - New TPD experiments on lunar simulants and samples initiated at GSFC (Jason McLain)

# Migration calculations

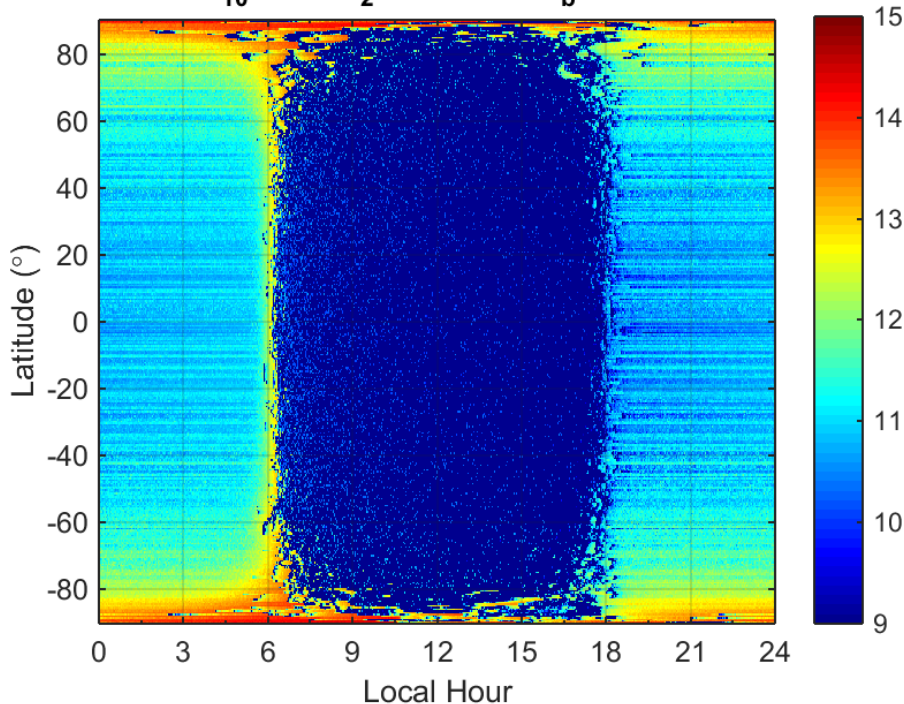


Adopted 23 gridded Diviner temperature maps,  
Resolution 0.5x0.5 deg

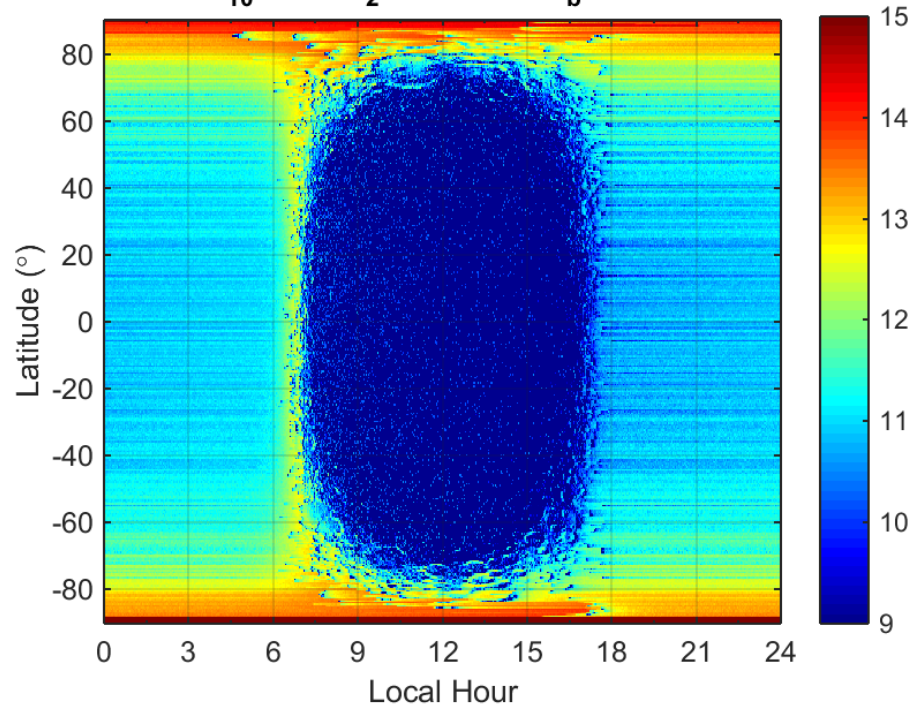
- \* Assume full vaporization of impactors (no contribution from trapped carbon in soil)
- \* Half CO, Half CO<sub>2</sub>
- \*  $\tau(\text{CO}_2) = 4.95\text{E}5 \text{ s}$ ,  $\tau(\text{CO}) = 1.3\text{E}6 \text{ s}$  (quiet Sun, charge exchange and electron impact dissociation/ionization losses not accounted yet)
- \*  $T_{\text{iv}} = 2000 \text{ K}$  here
- \* Recycled vapors assumed to stick
- \* Activation energy for desorption treated as free parameter,  $E_b = 0.5$  and  $E_b = 0.8 \text{ eV}$  (same for all soils)
- \*  $\text{Pre\_exponential} = 1.\text{E}13 \text{ /s}$

# Simulated CO<sub>2</sub> surface abundance

$\log_{10}(\text{adsCO}_2) \text{ (cm}^{-2}\text{)} \text{ for } E_b = 0.5 \text{ eV}$



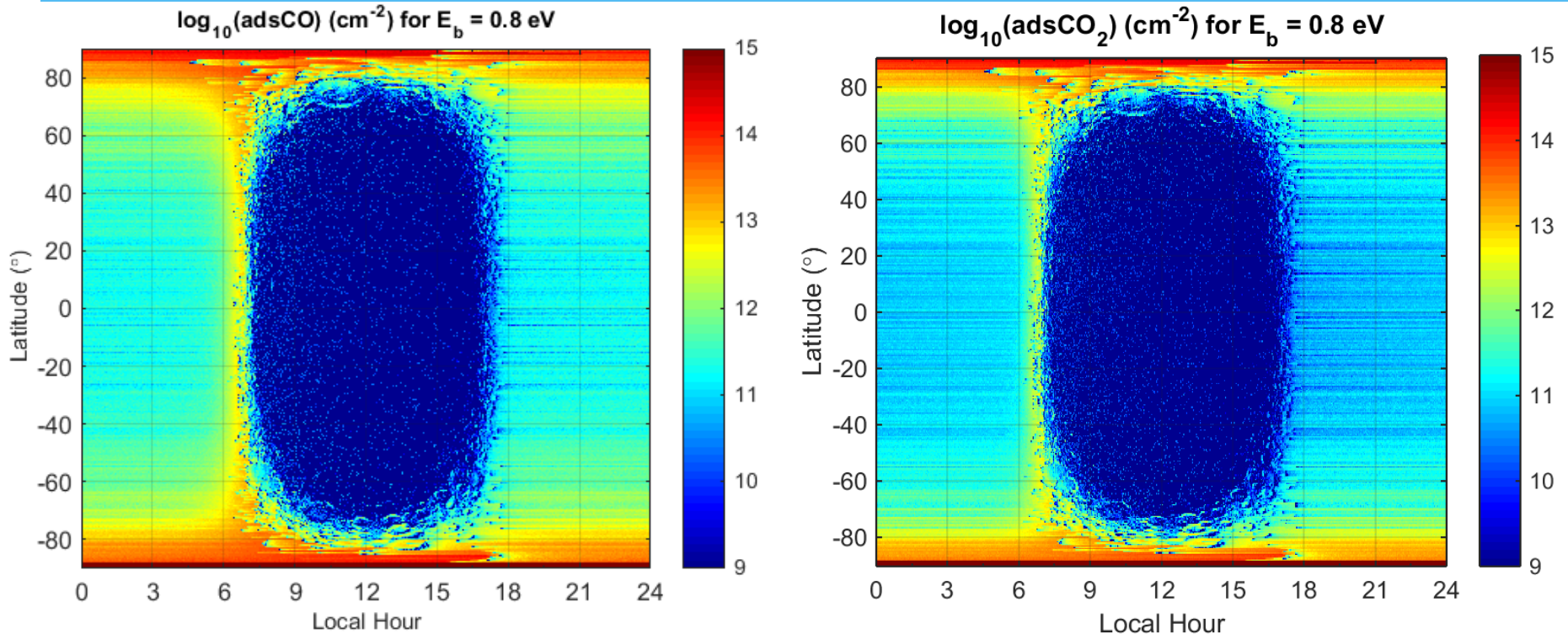
$\log_{10}(\text{adsCO}_2) \text{ (cm}^{-2}\text{)} \text{ for } E_b = 0.8 \text{ eV}$



- \* Coverage approaches or exceeds monolayer under such assumptions (must account for coverage dependence of  $E_{act}$ )
- \* Patchier polar volatile distribution as  $E_{act}$  decreases



# Surficial CO vs CO<sub>2</sub>

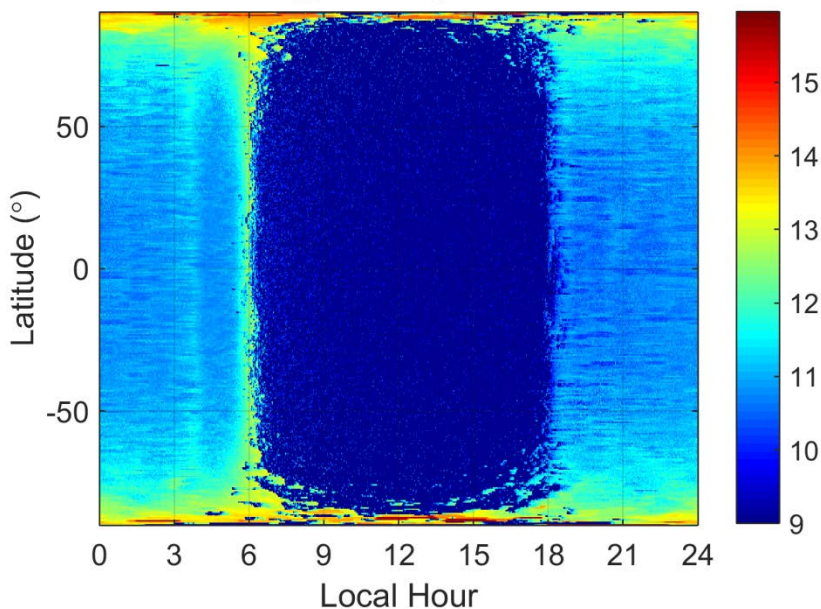


Assuming similar binding and source rates, we would expect significant amounts of CO to migrate to the poles as well (especially since its lifetime to dissociation is longer)

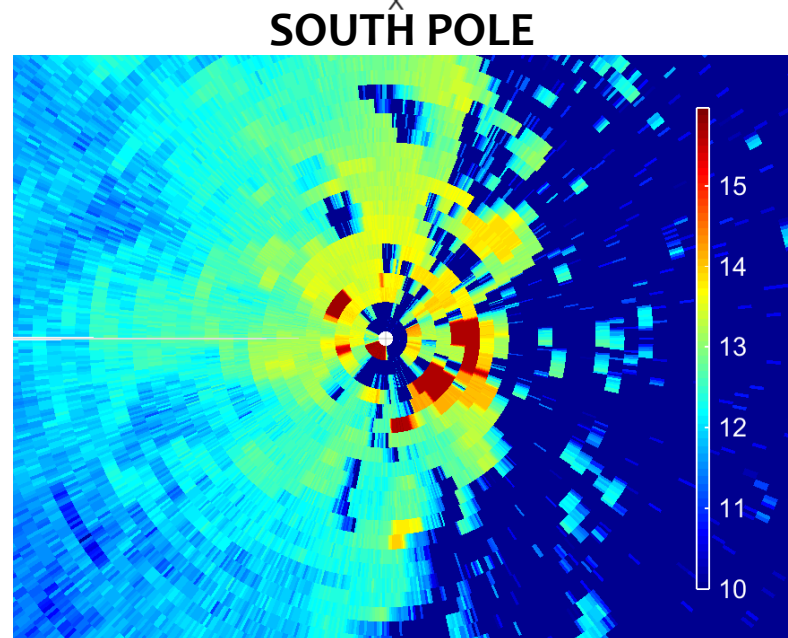
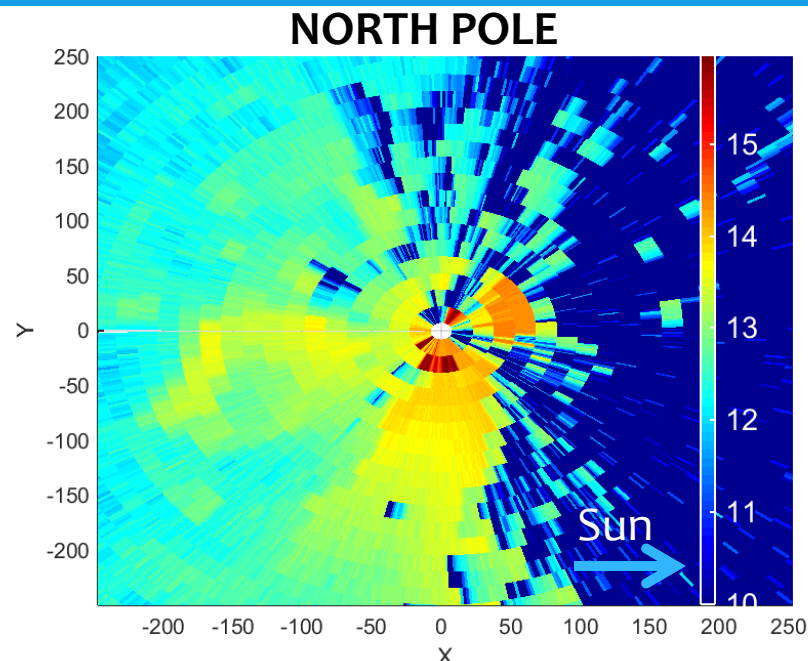
# Polar CO<sub>2</sub> frosts

$E_b = 0.5$  eV

$\log_{10}(\text{adsCO}_2) \text{ (cm}^{-2}\text{)}$



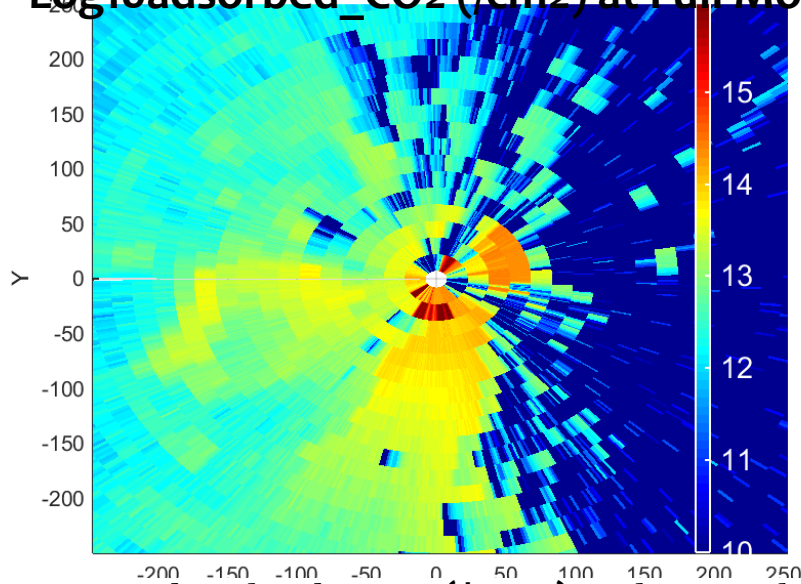
Time-dependent simulation run for 60 lunations  
Sinks: dissociation/ionization,  
adsorbate sputtering



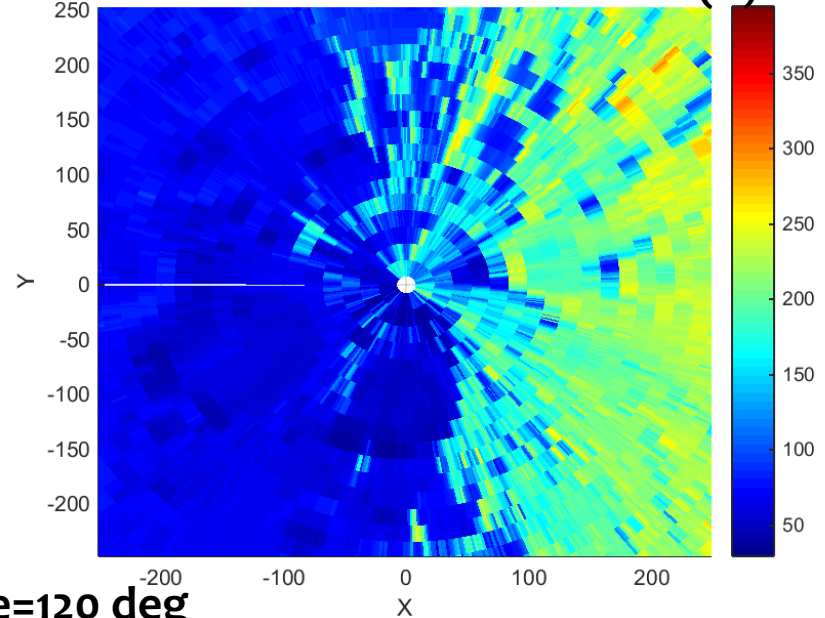


# Diurnal variation of polar CO<sub>2</sub> deposits

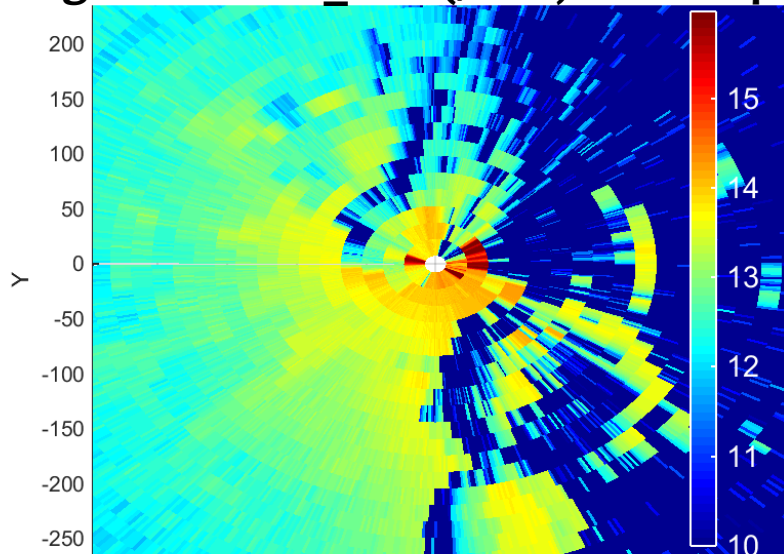
**Log10adsorbed\_CO2 (/cm2) at Full Moon**



**DIVINER TEMPERATURE (K)**



**Log10adsorbed\_CO2 (/cm2) at lunar phase=120 deg**



# Limits to carbon-bearing exospheric species prior to LADEE

## Detection methods:

- 1) Lunar Atmosphere Composition Experiment (LACE) NMS, ~25 deg North:  
possible detections of

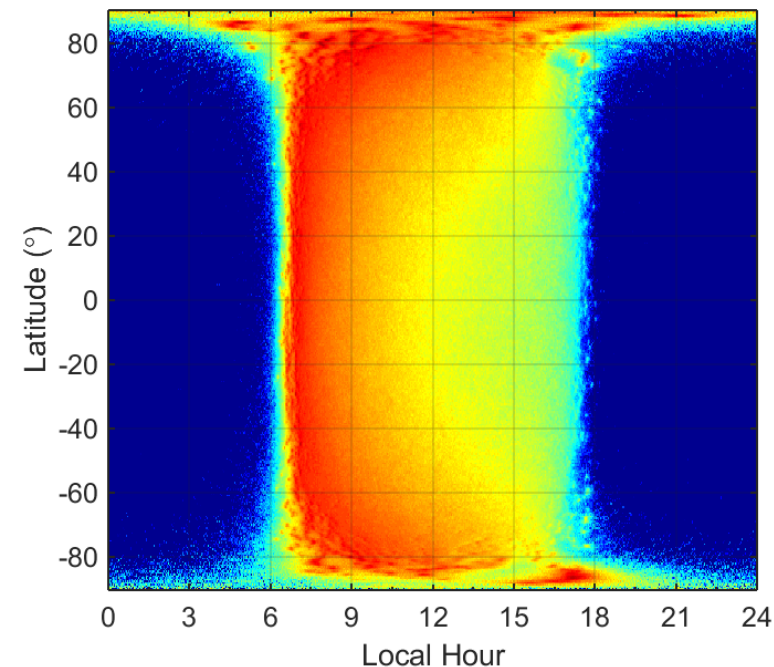
[CO]predawn=1.E3 cc, [CO<sub>2</sub>]predawn=1.E3 cc, [CH<sub>4</sub>]predawn=1.E4 cc

*“In all three cases, the signal increase began several hours before sunrise, and Continued to rise until the instrument became background saturated around sunrise”*

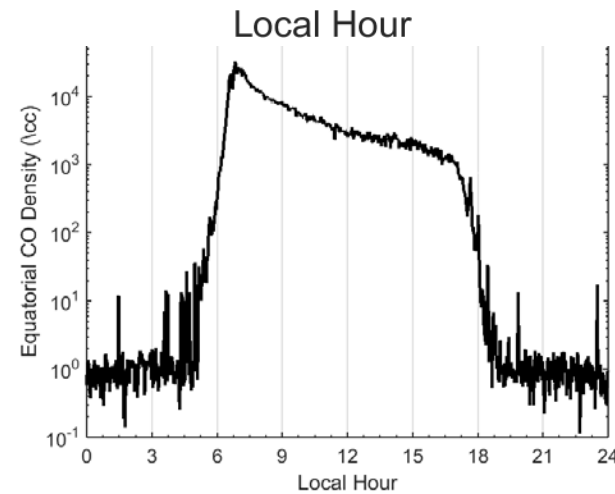
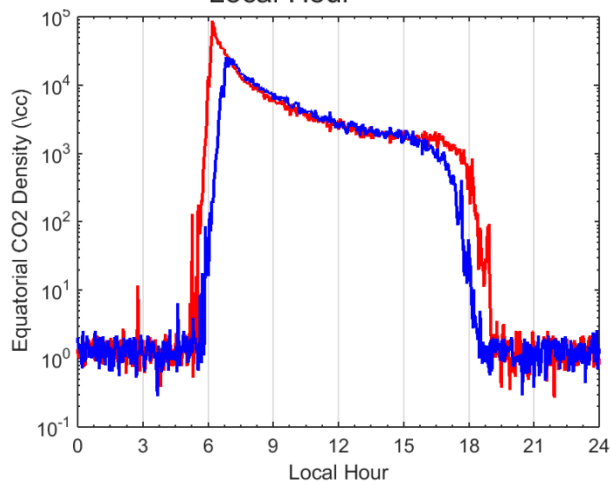
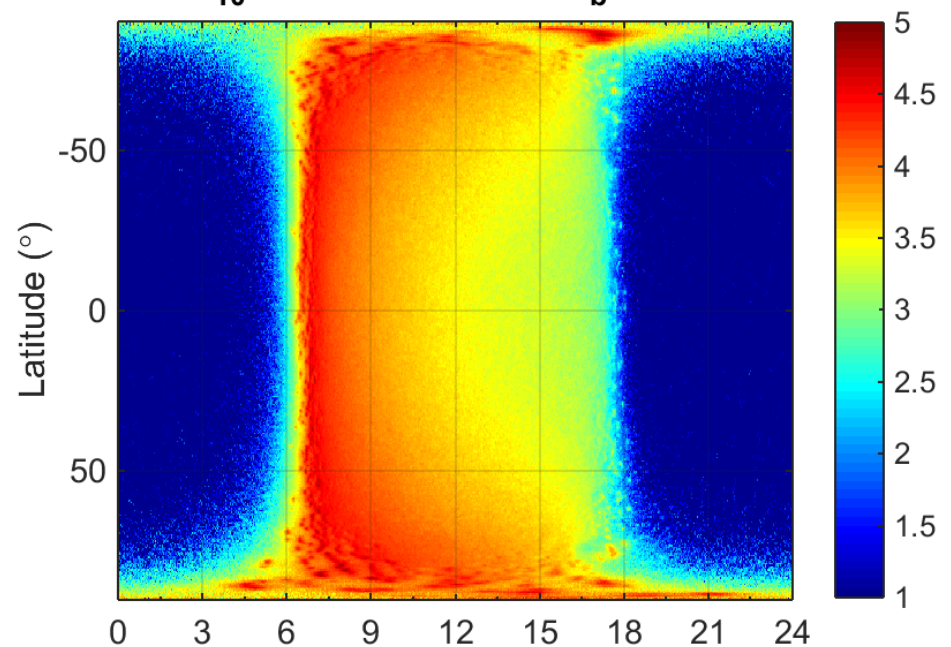
- 1) Apollo UVS: [CO<sub>g</sub>] < 14,000 cc (Feldman and Morrison, 1991)
- 2) LRO LAMP: [CO<sub>g</sub>] < 700 cc, polar nightside (Cook et al. 2013)

# Estimated Exospheric densities

$\log_{10}(\text{gasCO}_2) \text{ (cm}^{-2}\text{) for } E_b = 0.8 \text{ eV}$



$\log_{10}(\text{gasCO}) \text{ (cm}^{-2}\text{) for } E_b = 0.8 \text{ eV}$



# Conclusions

- \* Significant CO and CO<sub>2</sub> exospheric densities can be expected by the vaporization of micrometeoroids
- \* Estimates seemingly do not exceed known limits
- \* Can map global transport and deposition into poles at unprecedented 0.5x0.5 deg resolution
- \* Resulting polar deposits for CO and CO<sub>2</sub> would really be “frosts” (1-10 ML)