

RESIDUAL LOWER CONTINENTAL CRUSTAL COMPOSITIONS. S.R. Taylor, S.M. McLennan, R.J. Arculus and M.T. McCulloch, Research School of Earth Sciences, Australian National University, P.O. Box 4, Canberra, A.C.T. 2600, Australia.

The composition of the lower crust, which comprises 60-80% of the continental crust, remains a major unknown factor for models of terrestrial crustal evolution. The composition of the upper continental crust, exposed to weathering and erosion is well established as being of granodioritic composition<sup>1</sup>. For the lower crust, we lack those large scale natural sampling processes (such as production of clastic sediments or loess) which have simplified the task of arriving at upper crustal compositions<sup>2</sup>. Lower crustal samples are either random (as xenoliths in volcanic pipes) or from restricted outcrop areas of granulite terrains. The lower crust is almost certainly heterogeneous in detail, and may be further complicated by the presence of imbricate thrust sheets. One constraint is that the granodioritic (S.L.) rocks of the upper crust originated by partial melting within the crust, at depths of less than 40 km<sup>3,4</sup>. The lower crust must accordingly include many regions from which granitic melts (S.L.) have been extracted. If we can recognize such material in the scattered samples available it will provide valuable limitations on the bulk composition of the crust.

Various approaches are possible. One is to model the bulk crust, and calculate residual compositions following the extraction of granitic melts. In Table 1 (cols. D,E), such calculations are presented for the extraction of a minimum granitic melt from proposed total crustal compositions allowing for 10 and 20% extraction of melt. A slightly different approach is to extract the known upper crustal composition (Col. A) from the model total crust, assuming that the upper crust forms 33% (Col. F) and 20% (Col. G) of the total. These four compositions indicate that the lower crust should contain compositions high in Al<sub>2</sub>O<sub>3</sub>, CaO, low in K<sub>2</sub>O, with positive Eu anomalies (Eu/Eu\* > 1) and Nd<sub>v</sub>/Sm<sub>v</sub> ratios approaching chondritic values. Fig. 1 shows the upper crustal REE patterns, and the predicted REE patterns for the lower crust.

A second approach is to examine the composition of dry granulite samples, formed at lower crustal temperatures and pressures to see whether they match the model calculations. Fyfe<sup>4</sup> has pointed out the importance of removal of H<sub>2</sub>O and minimum granitic melts at the upper amphibolite grade of metamorphism, allowing the development of the anhydrous mineralogy typical of the granulite facies. Granulite facies rocks can be expected to show wide variations in composition due to several processes:

- (A) Development of granulite facies mineralogy in dry source rocks from which a granitic melt has been extracted during amphibolite facies metamorphism<sup>5</sup>.
- (B) Dehydration of source rocks with loss of an hydrous fluid phase, resulting in granulites depleted in alkalis and U<sup>6</sup>.
- (C) Dehydration without partial melting or loss of trace elements (eg Jeju complex, Brazil<sup>6</sup>).
- (D) Dehydration accompanied by loss of CO<sub>2</sub> (eg Southern India<sup>7</sup>).
- (E) Subsequent retrograde metamorphism to produce amphibolite facies mineralogy in which any or all of the above processes have operated.

Accordingly, much complexity is expected, and shown by the random examples of lower crustal compositions available. In Table 1, Cols. H to Q, data are given for a suite of Lewisian<sup>8</sup> and Scourian granulites<sup>9</sup>, granulite xenoliths from Lesotho<sup>10,11</sup> Bournac, France<sup>12</sup> and eclogites from Sauviat-sur-Vige, France<sup>13</sup>. These compositions are typified by high Al<sub>2</sub>O<sub>3</sub>, and CaO, low K<sub>2</sub>O and positive europium anomalies. Fig. 2 shows the REE patterns. Nd/Sm ratios (Table 1) are lower than either upper crust or total crustal estimates. The major and trace element compositions tend to show much variation, as noted by Nixon et al<sup>10</sup> in their study of the Lesotho xenoliths. In this example, minerals such as garnet show Eu enrichment and the development of the bulk rock REE pattern, with positive Eu anomalies, clearly predates the granulite facies metamorphism<sup>11</sup>. Accordingly, the extraction of granitic melts prior to granulite facies metamorphism<sup>4</sup> will change the bulk rock composition, including development of the Eu enrichment (since the granitic melts are typified by Eu depletion). Such compositional changes are observed, for example, in the Lewisian<sup>8</sup>. The well studied Scourian succession has been the subject of varying interpretations. Pride and Muecke<sup>9</sup> note the following arguments in favour of extraction of partial melts (a) anhydrous nature of the complex (b) incompatible element depletion (c) narrow range of mineral compositions (d) major element trends unlike those of upper crustal igneous rock sequences (e) REE abundances are lower than those typical of upper crustal rocks, with enrichment of europium. It is not possible to evaluate the relative frequency of such compositions, except to note that they are common.

#### References

- <sup>1</sup>Taylor, S.R. and McLennan, S.M. *Phil. Trans. Roy. Soc.* A301, p381 (1981).
- <sup>2</sup>Taylor, S.R., McLennan, S.M. and McCulloch, M.T. This volume.
- <sup>3</sup>Brown, G.C. and Fyfe, W.S., *Contr. Min. Pet.* 28, p.310 (1970).
- <sup>4</sup>Fyfe, W.S. *Phil. Trans. Roy. Soc.* A273, p.457 (1973).
- <sup>5</sup>Heier, K.S. *Phil. Trans. Roy. Soc.* A273, p.429 (1973).
- <sup>6</sup>Sighinolfi, G.P. et al., *Contr. Min. Pet.* 78, p.263 (1981).
- <sup>7</sup>Newton, R.C. et al., *Nature*, 288, p.45 (1980).
- <sup>8</sup>Drury, S.A. *Precambrian Res.* 7, p.237 (1978).
- <sup>9</sup>Pride, C. and Muecke, G.K. *Contr. Min. Pet.* 73, p.403 (1980).
- <sup>10</sup>Griffin, W.L. et al., in *The Mantle Sample* (Eds. F.R. Boyd and H.O.A. Meyer) Amer. Geophys. Union, p.59 (1979).
- <sup>11</sup>Rogers, N.W. and Hawkesworth, C.J. *Nature*, 299, p.409 (1982).
- <sup>12</sup>Dostal, J. et al., *E.P.S.L.* 50, p.31 (1980).
- <sup>13</sup>Bernard-Griffiths, J. and Jahn, B.M. *Lithos*, 14, p.263 (1981).
- <sup>14</sup>White, A.J.R. and Chappell, B.W. *Tectonophysics*, 43, p.7 (1977).

TAYLOR, S. R. et al.

Table 1. Lower crustal compositions and candidates

	A	B	C	D	E	F	G	H	I	J	K	L	M	N	O	P	Q
SiO <sub>2</sub>	66.0	58.0	76.5	53.4	55.9	54.0	56.0	61.3	50.6	52.2	52.6	57.5	66.5	65.3	49.5	48.5	41.9
TiO <sub>2</sub>	0.6	0.8	0.1	1.0	0.9	0.9	0.9	0.56	0.77	1.3	1.2	0.96	0.54	0.51	0.70	0.55	0.23
Al <sub>2</sub> O <sub>3</sub>	16.0	18.0	12.9	19.3	18.6	19.0	18.5	18.1	17.9	19.5	24.5	19.5	16.5	16.2	18.5	19.7	18.0
FeO <sub>T</sub>	4.5	7.5	1.1	9.1	8.2	9.0	8.3	5.4	8.8	9.2	5.9	5.7	4.0	4.2	6.6	5.5	13.3
MgO	2.3	3.5	0.1	4.4	3.9	4.1	3.8	2.4	8.2	5.9	2.0	2.8	2.0	2.3	8.3	8.7	27.6
CaO	3.5	7.5	0.7	9.2	8.3	9.5	8.5	5.1	9.4	8.7	7.9	6.6	5.4	5.3	14.5	14.8	5.3
Na <sub>2</sub> O	3.8	3.5	3.8	3.4	3.5	3.4	3.4	4.9	3.3	2.7	4.0	4.9	4.1	4.3	1.6	2.0	0.65
K <sub>2</sub> O	3.3	1.5	4.7	0.7	1.1	0.6	1.1	2.4	1.0	0.55	1.7	0.9	0.8	0.9	0.4	0.26	-
Eu/Eu*	0.64	1.0	0.45	1.35	1.12	1.22	1.10	1.97	2.4	1.46	3.14	1.69	1.29	1.34	1.31	1.50	1.70
Nd <sub>N</sub> /Sm <sub>N</sub>	1.88	1.41	1.72	1.14	1.29	1.09	1.26	1.79	1.24	-	-	-	-	-	1.29	1.26	1.24

A Upper Crust<sup>1</sup> B Total Crust<sup>1</sup> C Minimum melt composition<sup>14</sup> D Residue following 20% melt extraction from total crust  
 E Residue following 10% melt extraction<sup>14</sup> F Predicted lower crust composition after extracting 33% upper crust Col. 1 from Col. 2 G Predicted lower crust composition following extraction of 20% upper crust.  
 H Lewisian granulites, Ave of 8<sup>8</sup> I Lesotho garnet granulites PHN 1670, 2533, 2852 and L.13<sup>10,11</sup>  
 J Granulite xenolith, Bournac 3199<sup>12</sup> K Granulite xenolith, Bournac 3197<sup>12</sup> L Scourian granulite 65-18<sup>9</sup>  
 M Ibid, 67-34<sup>9</sup> N Ibid, 67-41<sup>9</sup> O Sauviat-Sur-Vige metagabbro 4736<sup>13</sup> P Ibid 4737 eclogite<sup>13</sup>  
 Q Ibid 4739 garnet peridotite<sup>13</sup>

