

EARLY EVOLUTION OF THE EARTH. Yutaka Abe and Takafumi Matsui, Geophysical Institute, Faculty of Science, University of Tokyo, Bunkyo-ku, Tokyo 113

If the earth was formed by accretion of planetesimals in the gas-free environment, the surface of a planet would melt due to the blanketing effect of an impact-induced atmosphere, as suggested by previous studies (1,2). Once melting occurs, the mass and composition of the proto-atmosphere will be significantly affected, because volatile materials such as  $H_2O$  and  $CO_2$  are redistributed between the atmosphere and magma ocean. The evolution of an impact generated  $H_2O$  atmosphere will be controlled by the solubility of water into silicate melt.

Generation of the proto-atmosphere depends on the surface temperature. (We assume the proto-atmosphere composed of only  $H_2O$ .) We need to consider three different source and sink mechanisms, depending on the temperature. i) When the surface temperature is lower than 900K, generation of  $H_2O$  atmosphere is mostly due to impact dehydration. (We can use an impact dehydration model and energy balance equation to calculate the surface temperature (2).) Once the peak impact pressure exceeds a critical value,  $P_{cr}$ , then some fraction,  $f$ , of  $H_2O$  retained in the planetesimals is added to the atmosphere, and so the mass of the atmosphere,  $M_a$ , is given by  $\dot{M}_a = fX_p\dot{M}$ , where  $X_p$  and  $\dot{M}$  are the  $H_2O$  content of planetesimals and mass accretion rate. ii) When the surface temperature exceeds 900K, serpentine becomes unstable and a dehydration reaction occurs without impact. Then the mass of the atmosphere  $M_a$  is given by  $\dot{M}_a = X_p\dot{M}$ . iii) When the surface temperature exceeds the melting temperature,  $M_a$  is controlled by the degree of partial melting,  $\alpha$ , at the surface and by the solubility of water into molten rock,  $X_{sol}$ .  $M_a$  is given by  $\dot{M}_a = (X_p - \alpha X_{sol})\dot{M}$ . Based on the experimental data compiled by Fricke and Reynolds (3),  $X_{sol}$  is given by  $X_{sol}(wt\%) = 2.08 \times 10^{-3} P^{0.54}$ , where  $P$  is the atmospheric pressure at the surface (in Pa). If there exists enough metallic iron to be able to react with  $H_2O$ , the  $H_2O$  in the atmosphere will decompose into  $H_2$  and FeO. Since the  $H_2/H_2O$  ratio of the atmosphere is a function of temperature, the evolution of an impact-induced  $H_2O$  atmosphere becomes very complicated. In the following, we follow the heterogeneous accretion model proposed by Matsui and Mizutani (4,5); that is, iron accretes first and silicate second.

Figure 1 shows the trace of the surface temperature for the standard model ( $f=0.2$ ,  $P_{cr}=228$ kbar,  $X_p=0.1\%$  by weight and accretion time=50m.y.). When the surface of the earth is molten, the surface temperature is nearly constant, due to the buffering function of the magma ocean. (When the degree of partial melting increases, the amount of water which dissolves into molten rock increases.) This leads to the decrease in the mass of the atmosphere, and decreases the surface temperature. Such a stable process keeps the surface temperature nearly constant. Figure 2 shows the growth curve of the  $H_2O$  atmosphere for the standard model. Generation of the atmosphere begins when the radius exceeds  $0.2R_0$ . Inflection of the growth curve at  $R=0.4R_0$  is due to the generation of magma ocean. It was shown that mass of the

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atmosphere is not determined by  $f$  and  $P_{cr}$  but determined by the accretion time and  $X_p$ . This is because of the buffering function of the magma ocean.  $P$  The final mass of the resulting atmosphere varies from  $3 \times 10^{20}$  kg to  $3 \times 10^{21}$  kg by changing the accretion time from  $5 \times 10^6$  to  $5 \times 10^7$  y and  $X_p$  from 0.1 to 1%. This value is very close to the mass of the present ocean ( $1.4 \times 10^{21}$  kg), and may imply that the mass of the present ocean was determined by the nature of magma ocean during earth's accretion.

## REFERENCES

- (1) Matsui, T. and Abe, Y. (1984) Lunar and Planetary Science XV, p.517-518. Lunar and Planetary Institute, Houston.
- (2) Abe, Y. and Matsui, T. (1985) Proc. Lunar Planet. Sci. Conf. 15th, in J. Geophys. Res., 90, in press.
- (3) Fricker, P.E. and Reynolds, R.T. (1968) Icarus, 9, 221-230.
- (4) Matsui, T. and Mizutani, H. (1977) Nature, 270, 506-507.
- (5) Matsui, T. and Schultz, P.H. (1984) Proc. Lunar Planet. Sci. Conf. 15th, in J. Geophys. Res., 89, c323-c328.

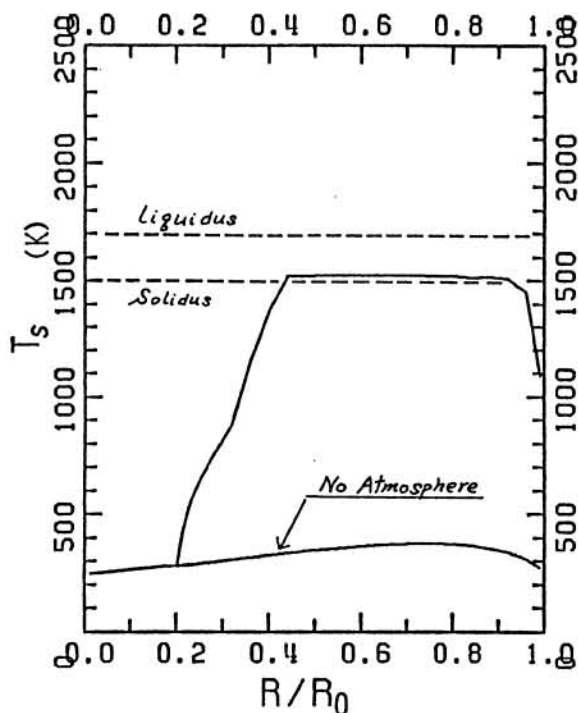


Figure 1

The surface temperature of the earth during accretion for the standard model.  $R_0$  is the radius of the present earth.

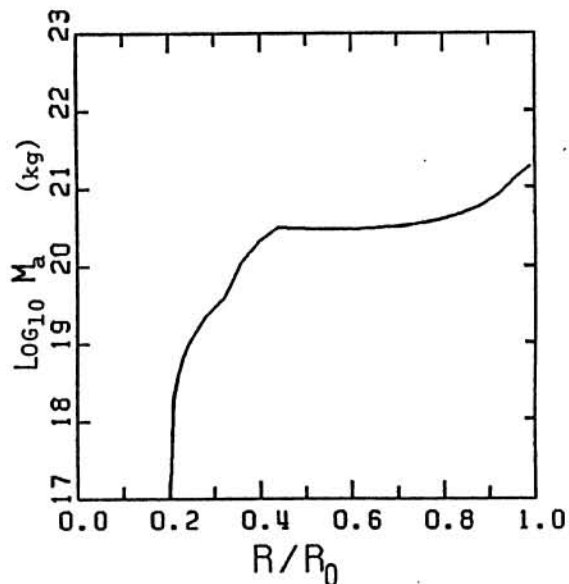


Figure 2

Growth curve of the  $H_2O$  atmosphere for the standard model.