

¹⁰Be PROFILES IN LUNAR SURFACE ROCK₃ 68815; K. Nishiizumi, M. Imamura¹, C. P. Kohl,² H. Nagai², K. Kobayashi³, K. Yoshida⁴, H. Yamashita⁵, R. C. Reedy⁶, M. Honda² and J. R. Arnold; Dept. of Chemistry, B-017, Univ. of Calif., San Diego, La Jolla, CA 92093 (U. S. A.), ¹Inst. for Nuclear Study, Univ. of Tokyo (Japan), ²Dept. of Chemistry, Nihon Univ., Tokyo (Japan), ³Res. Center for Nuclear Sci. and Tech., Univ. of Tokyo (Japan), ⁴Dept. of Chemistry, Univ. of Tokyo (Japan), ⁵Dept. of Physics, Univ. of Tokyo (Japan), ⁶Los Alamos National Lab, Los Alamos, NM 87545 (U. S. A.)

The comparison of ²⁶Al($t_{1/2} = 7.05 \times 10^5$ years) and ⁵³Mn(3.7×10^6 years) depth profiles in the surfaces of three lunar rocks (12002[1], 14321[2], and 68815[3]) with the theoretical SCR(solar cosmic rays) production profiles[4] indicates that the flux of solar protons over the past five to ten million years was similar to that during the past million years and that the average SCR spectrum and flux were characterized by a rigidity $R = 100$ MV and a flux $J = 70$ protons/cm² sec ($E > 10$ MeV, 4π). These calculations assume 0.5 - 2.2 mm/My erosion rates for the three rocks[3,5]. In this present work, we measured ¹⁰Be (1.6×10^6 years) in rock 68815 to understand the SCR production of this nuclide and to verify the SCR parameters.

Fourteen samples were separated from aliquant samples that we had used previously for ²⁶Al and ⁵³Mn measurements[3]. The samples measured were from three different zenith angles (A-45°, B-37°, and C-12°) and four different depths (0 - 4 mm). The ¹⁰Be measurements were carried out at the University of Tokyo tandem Van de Graaff accelerator[6]. The ¹⁰Be concentrations in samples from face C are shown in the figure. The ¹⁰Be results were adjusted to saturation using the ⁸¹Kr-Kr exposure age (2.04 ± 0.09 My)[7]. The ¹⁰Be profiles for face A and B are essentially the same as for face C, within error. The figure also shows data for two deeper samples. The ¹⁰Be profiles are extremely flat for all three phases and show a very slight increase with increasing depth. This shape is in contrast to the ²⁶Al and ⁵³Mn profiles, which show sharp increases in activity toward the surface due to SCR production of these nuclides[1,2,3].

It is necessary to subtract the GCR (galactic cosmic rays) produced ¹⁰Be from the observed ¹⁰Be to see the SCR component. The expected GCR production profile was calculated based on the Reedy-Arnold model[4] and is shown in the figure. The original model[4] and new cross sections[8] were used for both GCR and SCR calculations. The Reedy-Arnold GCR profile fits the 68815 data very well. However, there is no sign of SCR produced ¹⁰Be in this rock. The Reedy-Arnold SCR model predicts an SCR production of about 2 dpm ¹⁰Be/kg₂ in the surface layer using a flux with $R = 100$ MV and $J(>10 \text{ MeV}) = 70$ p/cm² s, the values we had obtained from ²⁶Al and ⁵³Mn. There are several possible explanations for the discrepancy between the model and the data: (1) The ¹⁰Be proton induced cross sections are too high, especially at the low energies that dominate the SCR flux. There are no cross section measurements below 135 MeV for protons on oxygen, which is the main target element for the production ¹⁰Be in our sample. Low energy cross sections for ¹⁰Be production from oxygen and other elements should be measured by AMS (accelerator mass spectrometry). (2) The average SCR flux and mean rigidity over the last two million years differed from the above values. In particular, the R_0 was lower than 100 MV. The Reedy-Arnold model shows that the ¹⁰Be production rates are very sensitive to changes in R_0 . Although the production rates of ²⁶Al and ⁵³Mn are less sensitive at near surface depths to changes in R_0 , a higher SCR flux and larger

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erosion rates would be required to fit those profiles if we use a lower R_0 . It is known that $R_0 = 100\text{MV}$ and $J = 70 \text{ p/cm}^2 \text{ s}$ are not unique parameters to fit the ²⁶Al and ⁵³Mn profiles in lunar surface rocks[5]. However, lowering the R_0 conflicts with the argument by Bhandari et al.[9] who proposed a higher R_0 than 100 MV. Reedy [10], using ⁸¹Kr data of Yaniv et al.[11], also found a somewhat higher R_0 for the period 3×10^5 years, but this is not necessarily a contradiction. (3) The calculated GCR production rates are too high near the surface due to the difficulty of estimating the Reedy-Arnold fluxes. This requires that the GCR production profile decrease about 20% from 1 g/cm^2 to the surface to obtain 2 dpm ¹⁰Be/kg SCR production in this region. There is no theoretical and experimental support for such an abrupt change.

Previous ¹⁰Be measurements in 12002, 14310, and 14321 using decay counting techniques[1,2] are in good agreement with the ¹⁰Be activity found in 68815 by AMS measurements and show no increase at the surface. Since substantially all the ¹⁰Be was produced by GCR, we can conclude that no significant changes in the GCR flux were observed during the last few million years.

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