

**CHRONOLOGY OF LUNAR GRANITE 12033,576: RESETTING OF Rb-Sr AND K-Ca ISOCHRONS;** C.-Y. Shih, H. Wiesmann and D. H. Garrison, Lockheed Engineering and Science Co., 2400 NASA Road 1, Houston, TX 77058; L. E. Nyquist and D. D. Bogard, NASA Johnson Space Center, Houston, TX 77058.

Lunar granite 12033,576 is a subsample of the "large" (~1 g) felsite 12033,507 which was identified from a collection of 4-10 mm particles from the 12033 soil sampled from the north rim of Head Crater in the eastern part of Oceanus Procellarum. Discordant ages of ~3.6, ~0.8, ~3.9 and ~2.2 Ga for this lunar granite were obtained, respectively, by the K-Ca, <sup>39</sup>Ar-<sup>40</sup>Ar and U-Pb zircon methods in previous studies and by the Rb-Sr method in this study. Assuming the granite crystallized ~3.9 Ga ago (zircon age), and was shocked by meteoritic impacts at 0.8 Ga ago (<sup>39</sup>Ar-<sup>40</sup>Ar age), the intermediate apparent ages by the Rb-Sr and K-Ca methods can be interpreted as reset by diffusion of the parent and daughter nuclides. The Rb-Sr age is less resistant to resetting than the K-Ca age, but more resistant than the <sup>39</sup>Ar-<sup>40</sup>Ar age.

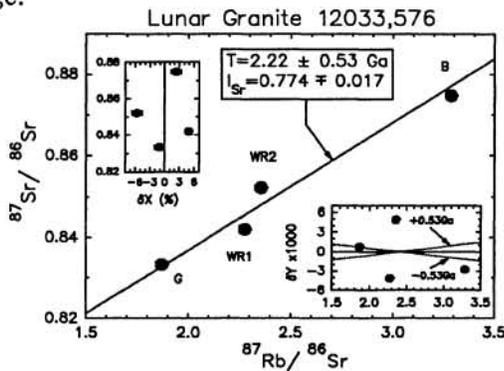


Figure 1. Rb-Sr mineral isochron of granite 12033,576.

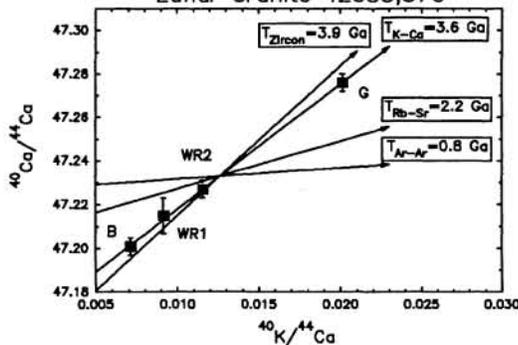


Figure 2. Discordant ages for granite 12033,576, shown on the K-Ca isochron diagram.

Ca isochron age and probably close to the U-Pb zircon age of ~3.9 Ga, but not older than the K-Ca model age of ~4.2 Ga.

**Diffusion model for resetting internal isochrons:** Discordant ages for the granite are summarized in a K-Ca isochron diagram (Fig. 2). Assuming that the granite crystallized at ~3.9 Ga and underwent shock metamorphism at 800 Ma, fractional exchanges (loss or gain), F(Sr) and F(Ca), of <sup>87</sup>Sr and <sup>40</sup>Ca, respectively, by diffusion in partially reset isotopic systems can be defined as  $F = (R_{3.9} - R_t) / (R_{3.9} - R_{0.8})$  where R is <sup>87</sup>Sr/<sup>86</sup>Sr or <sup>40</sup>Ca/<sup>44</sup>Ca for points on isochrons representing ages of 3.9 Ga, 0.8 Ga and the observed age t at a given parent/daughter ratio. Thus defined, F is given by the relative degree of isochron rotation, and can be calculated by a lever rule [7]. The calculated F(Sr) and F(Ca) for mineral separates B

**Rb-Sr mineral isochron for granite**

**12033,576:** Rb and Sr isotopic data for this subsample of large felsite 12033,507 [1] were obtained with procedures described in [2]. Two bulk samples, WR1 and WR2, and two mineral separates, G and B, handpicked from the whole rock sample crushed to <149 μm were analysed. G is a gray feldspar-rich sample and B is a dark sample containing brownish glasses of possible shock origin [1] as well as mafic minerals. K-Ca isotopic data for these samples yield a well-defined isochron age of 3.62±0.11 Ga [3]. Rb-Sr isotopic results are reported here for the same samples, and do not define a good isochron (Fig. 1). The best-fit line for the disturbed Rb-Sr system yields a young and imprecise age of 2.22±0.53 Ga (2σ) and a high initial <sup>87</sup>Sr/<sup>86</sup>Sr of 0.774±0.017 (2σ) using the York program [4]. Both K-Ca and Rb-Sr ages are younger than the upper concordia intersection age of 3.898±0.010 Ga obtained from the U-Pb data for zircons [5]. All these ages are considerably older than the <sup>39</sup>Ar-<sup>40</sup>Ar age of 800±15 Ma, interpreted as the age of Copernicus, determined from an aliquot of the bulk sample [6]. The high initial <sup>87</sup>Sr/<sup>86</sup>Sr and <sup>40</sup>Ca/<sup>44</sup>Ca values for the granite indicate that both the Rb-Sr and K-Ca isochrons could have been partially reset by the Copernicus impact. If so, the crystallization age for this granite could be significantly older than the K-

## AGE OF LUNAR GRANITE 12033,576: Shih C.-Y. et al.

and  $G$  are 0.58-0.60 and 0.18-0.16, respectively. These values can be used to calculate the diffusion parameter  $\alpha=(Dt/a^2)^{1/2}$  for Sr and Ca, using procedures analogous to those developed by [8] for the diffusion of Ar. The results for  $\alpha(\text{Ar})$ ,  $\alpha(\text{Sr})$  and  $\alpha(\text{Ca})$  are  $>0.8$ ,  $\sim 0.2$  and  $\sim 0.05$ , respectively, as presented in Fig. 3. The three solid lines represent the anticorrelation between diffusion time and diffusion coefficient for the three isotopic systems: K-Ar, Rb-Sr and K-Ca. Petrographic evidence suggests that the granite had experienced a shock-related thermal event at  $\sim 700^\circ\text{C}$  [1]. The  $D/a^2$  value for the K-Ar system at this temperature is determined from an Arrhenius diagram of temperature release  $^{39}\text{Ar}$  data to be  $\sim 10^{-6} \text{ sec}^{-1}$ . This temperature and  $D/a^2$  value correspond to an Ar diffusion time of several days needed to totally reset the K-Ar chronometer. The Rb-Sr and K-Ca isotopic data indicate that diffusion in the Rb-Sr and K-Ca system is, respectively,  $\sim 10\text{x}$  and  $\sim 400\text{x}$  slower than in the K-Ar system. These differences are in good agreement with the relative values of experimentally determined diffusion coefficients for Sr and Ca in granitic melts recommended by [9]. At  $700^\circ\text{C}$ , it would have required months to totally reset the Rb-Sr system in the granite, and years to totally reset the K-Ca system.

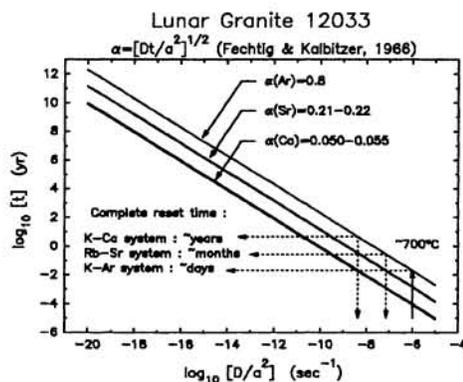


Figure 3. Diffusion time ( $t$ ) vs. parameter ( $D/a^2$ ).

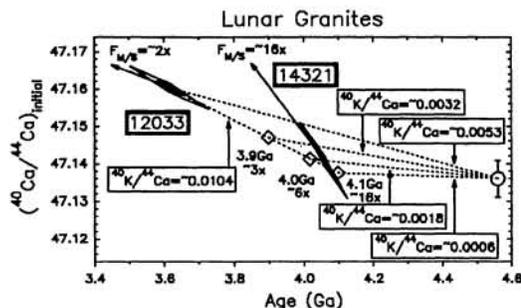


Figure 4. Age vs.  $I(\text{Ca})$  for lunar granites.

### Crystallization age and petrogenetic implications:

Ages and initial  $^{40}\text{Ca}/^{44}\text{Ca}$  ratios of lunar granites 14321 and 12033 [3,10] are represented by error parallelograms in Fig. 4. Dotted lines represent  $^{40}\text{K}/^{44}\text{Ca}$  growth curves. For a simple two-stage model, the  $^{40}\text{K}/^{44}\text{Ca}$  ratio in the source of granite 14321, which has concordant ages [10,11], was calculated to be  $\sim 0.0018$ , similar to values found for quartz monzodiorites [12-14], suggesting large K/Ca fractionations ( $F_{\text{M/S}} \sim 16\text{x}$ ) during granite formation. Large enrichments of K/Ca can be produced by silicate liquid immiscibility (SLI) processes [e.g. 15,16]. The  $^{40}\text{K}/^{44}\text{Ca}$  ratio calculated for the source of granite 12033 formed at the K-Ca isochron age of 3.6 Ga is so high that it either must have been derived from an already granitic source, or have assimilated a large amount ( $\sim 40\%$ ) of ancient granitic materials prior to its crystallization. Assuming a similarly large  $F_{\text{M/S}} \sim 16\text{x}$  for this granite as for 14321, the 12033 granite could have been derived from a low K/Ca source similar to that for granite 14321  $\sim 4.1$  Ga ago in a plutonic environment perhaps associated with the parental magma of some Mg-suite rocks.

### References:

- [1] Warren P. H. et al., (1987) PLPSC 17th, E303-E313.
- [2] Nyquist L. E. et al. (1990) GCA 54, 2195-2206.
- [3] Shih C.-Y et al. (1992) LPS XXIII, 1289-1290.
- [4] York D. (1966) Can. J. Phys. 44, 1079-1086.
- [5] Meyer C. Jr. et al., (1989) LPS XX, 691-692.
- [6] Bogard D. D. et al. (1992) LPS XXIII, 133-134.
- [7] Nyquist L. E. et al. (1979) 43, 1057-1074.
- [8] Fechtig H. and Kalbitzer S. (1966), In Potassium Argon Dating (ed. O.A. Schaeffer and J. Zaehring), Springer-Verlag New York Inc., 66-106.
- [9] Jambon A. (1982) JGR 87, B-13, 10797-10810.
- [10] Shih C.-Y et al. (1992) GCA (submitted).
- [11] Shih C.-Y et al. (1985) GCA, 49, 411-426.
- [12] Ryder G. (1976) EPSL 29, 255-268.
- [13] Marvin U.B. et al. (1991) PLPSC 21st, 119-135.
- [14] Jolliff B. L. (1991) PLPSC 21st, 101-118.
- [15] Watson E. B. (1976) Contrib. Mineral. Petrol. 56, 119-134.
- [16] Longhi J. (1990) PLPSC 20th, 13-24.