

STRUCTURE AND RHEOLOGY OF PARTIALLY MOLTEN AMMONIA-WATER ICES ; D.L. Goldsby and D.L. Kohlstedt (Dept. of Geology and Geophysics, Univ. of Minnesota, Mpls, MN, 55455)

The physical properties of partially molten planetary materials, including permeability and rheology, are fundamentally dependent on the grain-scale distribution of the melt phase [1,2]. The results of high-pressure creep experiments on fine-grained partially molten ammonia-water ices reported in [3] have been analyzed in the context of recent observations of the wetting characteristics of the melt in two-phase ammonia-water ices. These creep experiments explored the effects of temperature T , strain rate $\dot{\epsilon}$, melt fraction ϕ , and grain size on creep strength. Experiments were conducted in a gas-medium apparatus designed for cryogenic use under the following conditions: $3.5 \times 10^{-6} < \dot{\epsilon} < 3.5 \times 10^{-4} \text{ s}^{-1}$, $160 < T < 220 \text{ K}$, and confining pressure $P_c = 50 \text{ MPa}$. Samples were prepared containing 1, 5 or 8 wt% ammonia, yielding calculated melt fractions, respectively, of 3-4, 15-20, and 24-32 vol% from the peritectic temperature $T_p = 176 \text{ K}$ to $T = 220 \text{ K}$. Partial melting of samples containing 1 wt% ammonia resulted in less than a factor of 2 decrease in strength over the entire experimental temperature range. Samples containing 5 and 8 wt% ammonia, however, exhibit a more dramatic loss of strength -- up to a factor of 5 -- at temperatures just above the peritectic temperature. Scanning electron microscope (SEM) micrographs of samples equilibrated at $T = 185 \text{ K}$ for 24 h at a confining pressure $P_c = 1 \text{ atm}$ reveal that at low melt fractions (e.g., 3 vol%) melt is confined to triple junctions, that is, grain faces remain melt-free. In contrast, at higher melt fractions, ammonia-water melt completely surrounds Ice I grains, so that the rheologically critical melt fraction ϕ_c is exceeded, as shown in Figure 1.

The presence of melt in a polycrystalline aggregate can have a profound effect on rheology, both in the dislocation and in the diffusion creep regimes [4,5]. Melt increases the rate of dislocation creep due to the stress enhancement caused by a decrease in load-bearing area. Likewise, melt increases the rate of diffusional creep due not only to an enhancement in stress but also to the presence of high diffusivity melt-filled pathways [6]. Thus, the distribution of the melt phase, often characterized by a dihedral or wetting angle θ , governs the creep strength of a partially molten aggregate. For $0^\circ < \theta < 60^\circ$, at low melt fractions, melt forms an interconnected network along three-grain and through four-grain junctions; at high melt fractions, the grains become completely surrounded by melt [1].

Previous microstructural analysis of ammonia-water ices deformed at high pressure has relied on replication of the outer sample surface by the indium jackets used to seal the samples [7]. This technique, however, does not provide adequate spatial resolution to measure the dihedral angle. A cryogenic temperature stage was developed for an environmental scanning electron microscope (ESEM) in the High Resolution Microscopy Center at the University of Minnesota that allows *in situ* microstructural characterization of ammonia-water ice samples. This cryosystem consists of a copper sample stage coupled to a self-pressurized liquid nitrogen (LN_2) dewar; subsolidus sample temperatures ($< T = 176 \text{ K}$) are maintained by circulating LN_2 through the sample stage at a controlled rate. The system provides stable temperature control in the range $100 \text{ K} < T < 273 \text{ K}$. The ESEM permits samples to be analyzed at much higher chamber pressures than is possible in a conventional SEM (i.e. $\sim 10^{-6} \text{ Torr}$ vs. $\sim 10^{-5} \text{ Torr}$). In addition, sputter-coating of a high-conductivity material on the ice surface is not necessary in the ESEM, as would be required in a conventional SEM. The higher chamber pressures in the ESEM, combined with the lower temperature environment, reduces the rate of sample

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sublimation to a negligible level. Condensation on the sample surface is minimized through the use of a dry imaging gas, such as N_2 , instead of water vapor; special sample exchange techniques are also used to minimize condensation.

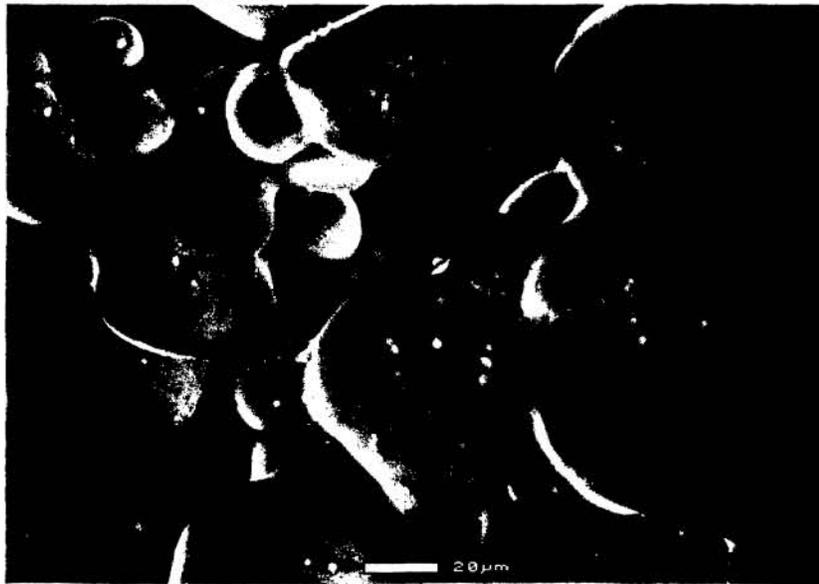


Figure 1- ESEM micrograph of 8 wt% NH_3 sample equilibrated for $t=24$ h at 1 atm pressure, $T=185$ K. Calculated melt fraction $\phi \approx 0.24$. Note that Ice I grains are completely separated by melt.

With this newly developed capability, SEM observations have been made in the ammonia-water system. At $T=185$ K for the composition $X_{NH_3}=0.01$, the dihedral angle is quite small, with $\theta \leq 20^\circ$. Hence, in the ammonia-water system, just above the peritectic temperature and for low melt fractions, melt forms a three-dimensional, interconnected network. At higher melt fractions, as in the $X_{NH_3}=0.05$ and 0.08 cases, Ice I grains are completely separated by melt, as shown in Figure 1. These SEM observations, coupled with the results of the creep experiments performed at hypersolidus temperatures for the samples with 5 and 8 wt% ammonia demonstrate that ϕ_c must be substantially less than 15 vol%. Consequently, partially molten ammonia-water ice near the surface and further within icy planetary bodies will undergo a dramatic loss of strength at relatively small melt fractions (i.e., at relatively water-rich compositions). These results also imply that the strength of partially molten ammonia-water ices will be critically dependent on the permeability, which will control the amount of melt retained in the ice body [8,9].

References

- [1] Bulau J.R. et al. (1979), *JGR*, 84, 6102-6108.
- [2] Kohlstedt D.L. (1992), in *AGU Geophys. Monogr.* 71, 103-121.
- [3] Goldsby D.L. (1993), *Proceedings of 24th LPSC Meeting*, 543-544.
- [4] Beeman M.L. and Kohlstedt D.L. (1993), *JGR*, 98, 6443-6452.
- [5] Chopra P.N. and Kohlstedt D.L. (1994), in *Magmatic Systems*, *in press*.
- [6] Cooper R.F. et al. (1989), *Acta Met.*, 37, 1759-1771.
- [7] Durham W.B. (1993) *JGR*, 98, 17667-17682.
- [8] Riley G.N. et al. (1990), *Geophys. Res. Lett.*, 17, 2101-2104.
- [9] Daines M.J. and Kohlstedt D.L. (1993), *Geophys. Res. Lett.*, *in press*.