

REEVALUATION OF CALCIUM-41 IN CM AND CV REFRACTORY INCLUSIONS Ming-Chang Liu¹, Kevin D. McKeegan¹, Andrew M. Davis² and Trevor R. Ireland³ ¹Department of Earth and Space Sciences, UCLA, Los Angeles, CA 90095 (mcliu@ess.ucla.edu), ²Department of the Geophysical Sciences, Enrico Fermi Institute, and Chicago Center for Cosmochemistry, University of Chicago, Chicago, IL 60637, ³Research School of Earth Sciences, ANU, Canberra, ACT, Australia.

Introduction: ⁴¹Ca is of importance in early solar system chronology as it provides a strict constraint regarding the time span between nucleosynthesis and preservation in refractory solids due to its extremely short half-life (~0.1 My). Its presence is detected through excesses of radiogenic ⁴¹K as first demonstrated by Srinivasan and colleagues [1,2] in Efremovka CAIs. Using a small geometry CAMECA 4f ion microprobe, these workers found that the degree of ⁴¹K excess correlated with the ⁴⁰Ca/³⁹K ratio, yielding an initial ⁴¹Ca/⁴⁰Ca~1.4×10⁻⁸. The existence of ⁴¹Ca in the solar system, if it came from a stellar source, implies a very short timescale for transit and incorporation. On the other hand, ⁴¹Ca can also be made locally by irradiation, although the physical conditions that produce ⁴¹Ca~1.4×10⁻⁸×⁴⁰Ca would lead to underproduction of ²⁶Al by a factor of 10 compared to the canonical level [3]. Interestingly, ⁴¹Ca was soon discovered to be correlated with the presence or absence of ²⁶Al in CM hibonite grains [4], implying ²⁶Al and ⁴¹Ca were derived from a common stellar origin. The absence of short-lived ⁴¹Ca and ²⁶Al in some platy hibonite crystals has been interpreted as due to a late injection of these isotopes into the solar nebula, after formation of the hibonite [4–7].

Based on the existing correlation of ⁴¹Ca with ²⁶Al found by [4,6], McKeegan et al. [8] suggested that the initial ⁴¹Ca abundance in the early solar system could have been significantly higher (potentially up to ~4×10⁻⁷×⁴⁰Ca) if the true initial ²⁶Al abundance of the solar nebula was “supracanonical” [9].

Despite its importance for understanding early solar system chronology, the ⁴¹Ca abundance in CAIs has never been quantitatively confirmed at other laboratories. Promising results were obtained with the SHRIMP-RG ion microprobe at Stanford University [10], however systematic uncertainties in the mass spectrometry (corrections for doubly ionized species and for peak-tailing) precluded a definitive answer regarding the level of ⁴¹Ca in E44. Here we report results of potassium isotope measurements in CM hibonite and the E44 CAI using the UCLA CAMECA IMS 1270 ion probe and we confirm the presence of ⁴¹Ca in early solar system materials.

Experimental: The isotopic compositions of potassium in both CM hibonite (6 PLATy hibonite Crystals “PLACs”, 2 Spinel-HIBonite spherules “SHIBs” and one Blue AGgregate “BAG”) and a piece of E44 CAI

were determined by adopting the procedure of [2] with some modifications. Polished, epoxy-mounted samples were sputtered with a 22.5 KeV ¹⁶O⁻ primary beam with an intensity of ~4 nA to generate sufficient count rates on Ca and K isotopes. Secondary ions were collected in a peak-jumping mode under the mass resolution (M/ΔM) of ~7000, which enables us to well separate ⁴⁰CaH from ⁴¹K as well as some minor interferences from the peaks of interest. However, another major interference at mass 41, (⁴⁰Ca⁴²Ca)⁺⁺, was still unresolvable from ⁴¹K so that its magnitude had to be assessed indirectly. For this, we followed the approach from [11] and [2], assuming the equality holds:

$$\frac{({}^{40}\text{Ca}^{42}\text{Ca})^{++}}{{}^{42}\text{Ca}^+} = \frac{({}^{40}\text{Ca}^{43}\text{Ca})^{++}}{{}^{43}\text{Ca}^+},$$

where the (⁴⁰Ca⁴³Ca)⁺⁺ signal can found at m/e=41.5. In most cases, this peak was hard to locate due to its extremely low intensity (<1 cps); therefore, its position was determined by setting a reference peak of (⁴⁰Ca²⁷Al¹⁶O)⁺⁺, which is 9 mass heavier than (⁴⁰Ca⁴³Ca)⁺⁺. It is extremely important that this correction be made accurately since the contribution of (⁴⁰Ca⁴²Ca)⁺⁺ comprises 1–80% of total ⁴¹K signal depending on K abundance. In addition to (⁴⁰Ca⁴²Ca)⁺⁺, special efforts were also made to evaluate the contribution of scattered ions from ⁴⁰Ca⁺ and ⁴⁰CaH⁺ tails at mass 41. The former was monitored by setting a series of analyses at the masses between ⁴⁰Ca and ⁴¹K. Scattered ⁴⁰Ca ions on ⁴¹K were then estimated at the level of ~1×10⁻¹⁰×⁴⁰Ca by extrapolating signals to mass 41. The latter was assessed by obtaining the count rate at mass 41.95 and applying the following relationship:

$$[{}^{40}\text{CaH}]_{\text{tail}} = \left[\frac{41.95}{42\text{Ca}^+} \right] \times [({}^{40}\text{CaH})^+]$$

This contribution of (⁴⁰CaH)_{tail} to ⁴¹K was found to be on the order of ~2.3×10⁻⁵×⁴⁰CaH and corrected for. In general, this correction is relatively minor (<1%) compared to that of (⁴⁰Ca⁴²Ca)⁺⁺ as ⁴⁰CaH signal was suppressed by the use of liquid nitrogen. The dynamic background of the counting system was also measured overnight when analyses were not performed. It was found within the range of 0.003 and 0.007 counts per second. The true ⁴¹K/³⁹K in a sample can be calculated after stripping off these interferences:

$$\left[\frac{{}^{41}\text{K}}{{}^{39}\text{K}} \right]_{\text{true}} = \left[\frac{{}^{41}\text{K}}{{}^{39}\text{K}} \right]_{\text{measured}} - \left[\frac{({}^{40}\text{Ca}^{43}\text{Ca})^{++}}{{}^{43}\text{Ca}} \times \frac{{}^{42}\text{Ca}}{{}^{39}\text{K}} \right] - \left[\frac{{}^{40}\text{CaH}}{39\text{K}} \right]_{\text{tail}}$$

Secondary ions were counted by the axial ion counter with the mass sequence of 38.8, ³⁹K, ⁴¹K,

$^{40}\text{CaH}^+$, $^{40}\text{Ca}^{43}\text{Ca}^{++}$, $^{40}\text{Ca}^{27}\text{Al}^{16}\text{O}^{++}$, ^{42}Ca and ^{43}Ca . In addition to 10–15 minute presputtering to minimize surface K contaminations, a field aperture was also introduced to filter out scattered K ions from the vicinity. Counting times on ^{39}K , ^{41}K and $^{40}\text{Ca}^{43}\text{Ca}^{++}$ were adjusted on the basis of their intensities for achieving enough counting statistics. ^{40}CaH , $^{40}\text{Ca}^{27}\text{Al}^{16}\text{O}^{++}$, ^{42}Ca and ^{43}Ca were typically counted for 10 seconds or less. Each measurement consists of 60–100 cycles; the total duration ranges from 1.5 to 3.5 hours. Charging corrections and mass calibration were performed every 5 cycles automatically to compensate the drifts during a long analysis. A series of terrestrial standards with a wide range of Ca/K ratios were analyzed to ensure the K mass spectrometry was properly functional. All measured $^{41}\text{K}/^{39}\text{K}$ ratios were normalized to the terrestrial reference ($^{41}\text{K}/^{39}\text{K}=0.072$). Even though ^{40}K was not measured, the finding of “normal” $^{41}\text{K}/^{39}\text{K}$ within errors throughout all the terrestrial standards was indicative of a small instrumental mass fractionation, which is negligible for the determination of ^{41}K excesses in extraterrestrial samples. In addition, the measured ratios of $^{42}\text{Ca}/^{43}\text{Ca}$ in standards were very close to the terrestrial value ($^{40}\text{Ca}/^{44}\text{Ca}=47.153$, $^{42}\text{Ca}/^{44}\text{Ca}=0.31221$, $^{43}\text{Ca}/^{44}\text{Ca}=0.06486$ [12]), implying a minor instrumental fractionation of Ca ($\sim 1\text{--}2\%$ amu) and insignificant Sr interferences at masses 42 and 43. Therefore, the true $^{40}\text{Ca}/^{39}\text{K}$ ratio of the measured phase is calculated with $^{40}\text{Ca}/^{39}\text{K} = ^{42}\text{Ca}/^{39}\text{K} \times ^{40}\text{Ca}/^{42}\text{Ca}$. A relative sensitivity factor of 2.28–2.31 between Ca and K in an ion microprobe was determined by comparison to the NBS 614 glass with known Ca and K concentrations (CaO=12 wt% and K=30 ppm).

Result and Discussion: All measured K isotopic ratios in the samples were subject to variable amount of corrections for $(^{40}\text{Ca}^{42}\text{Ca})^{++}$, ranging from 1 to 80%. Corrections for $^{40}\text{CaH}_{\text{tail}}$ were insignificant ($<0.1\%$) due to high mass resolution and the use of liquid nitrogen. CM hibonite, regardless of their morphological type, did not exhibit any resolvable ^{41}K excesses. In three PLAC samples (Mur-P2, Mur-P7 and Mur-P9), wherein K was extremely depleted ($^{40}\text{Ca}/^{39}\text{K}$ up to 1.2×10^7 , 4.8×10^7 and 1.7×10^7 , respectively), only an upper limit of $^{41}\text{K}/^{39}\text{K}$ ratios could be obtained. This indicated that the measured signal at mass 41 was all coming from $(^{40}\text{Ca}^{42}\text{Ca})^{++}$ and a fraction of $^{40}\text{CaH}_{\text{tail}}$. Our result in PLAC samples agrees with the study of [4]; whereas large analytical uncertainties in our measurements and the lack of sufficient fractionation of Ca from K in SHIBs prevented us from revealing any radiogenic potassium. On the other hand, large radiogenic ^{41}K excesses were found in two of the nine measurements in pyroxene of the E44 CAI; one of which showed a $\delta^{41}\text{K}$ enrichment of 8400% relative to the terrestrial refer-

ence and $^{40}\text{Ca}/^{39}\text{K}$ of 3.8×10^7 , yielding an inferred initial $^{41}\text{Ca}/^{40}\text{Ca} \sim 1.5 \times 10^{-8}$, which is in perfect agreement with the result from [2], and the other point indicated a model $^{41}\text{Ca}/^{40}\text{Ca}$ ratio of 7×10^{-9} (Fig 1). The remainders broadly overlapped the terrestrial value within errors, regardless of $^{40}\text{Ca}/^{39}\text{K}$ ratios, except for two points which could have possibly been due to overcorrection of $(^{40}\text{Ca}^{42}\text{Ca})^{++}$.

Our results show that PLACs and one BAG are devoid of both ^{41}Ca and ^{26}Al , indicating that these two isotopes were both lacking in the early solar system for a period of time, most likely prior to the formation of CAIs because of the presence of $\Delta^{26}\text{Mg}^*$ deficits in some of the PLACs [13]. This fits into the late injection model [5,7] well; however, the absence of ^{41}Ca in spinel-hibonite spherules, in which ^{26}Al is present at the canonical level, still prevents us from testing the relationship quantitatively.

On the contrary, $^{41}\text{Ca}/^{40}\text{Ca}$ ratio of 1.5×10^{-8} in the E44 CAI was confirmed in one pyroxene spot. We still have no evidence of a higher initial ^{41}Ca and more K-isotope studies in CAIs that reveal “supracanonical” ^{26}Al abundances are certainly needed. Nevertheless, if 1.5×10^{-8} is truly the solar system initial, the other well-resolved point, which yields a lower apparent $^{41}\text{Ca}/^{40}\text{Ca}$ of $\sim 0.7 \times 10^{-8}$, could have been due to isotopic resetting ~ 0.1 My after the CAI formed.

References: [1] Srinivasan et al. (1994) *ApJL*, **431**, 67–70 [2] Srinivasan et al. (1996) *GCA*, **60**, 1823–1835 [3] Lee et al. (1998) *ApJ*, **506**, 898–912 [4] Sahijpal et al. (1998) *Nature*, **391**, 559–562 [5] Sahijpal et al. (1998) *ApJL*, **509**, 137–140 [6] Sahijpal et al. (2000) *GCA*, **64**, 1989–2005 [7] Goswami et al. (2001) *ApJ*, **549**, 1151–1159 [8] McKeegan et al. (2004) *MAPS*, **39**, A5224 [9] Young et al. (2005) *Science*, **5719**, 223–227 [10] Ireland et al. (1999) *MAPS*, **34**, A57 [11] Hutcheon et al. (1984), *LPS*, **24**, 387–388 [12] Niederer et al. (1984) *GCA*, **48**, 1279–1293 [13] Liu et al. (2007) in prep.

