

MINERALOGY, PETROLOGY, CHRONOLOGY, AND EXPOSURE HISTORY OF THE CHELYABINSK METEORITE AND PARENT BODY. K. Righter¹, P. Abell¹, D. Agresti², E. L. Berger³, A.S. Burton¹, J.S. Delaney⁴, M.D. Fries¹, E.K. Gibson¹, R. Harrington⁵, G. F. Herzog⁴, L.P. Keller¹, D. Locke⁶, F. Lindsay⁴, T.J. McCoy⁷, R.V. Morris¹, K. Nagao⁸, K. Nakamura-Messenger¹, P.B. Niles¹, L. Nyquist¹, J. Park⁴, Z.X. Peng⁹, C.-Y. Shih¹⁰, J.I. Simon¹, C.C. Swisher, III⁴, M. Tappa¹¹, and B. Turrin⁴. ¹NASA-JSC, Houston, TX 77058; ²Department of Physics, University of Alabama at Birmingham, Birmingham, AL 35294-1170; ³GeoControl Systems Inc. – Jacobs JETS contract – NASA JSC; ⁴Rutgers Univ., Wright Labs-Chemistry Dept., Piscataway, NJ; ⁵UTAS – Jacobs JETS Contract, NASA-JSC; ⁶HX5 – Jacobs JETS Contract, NASA-JSC; ⁷Smithsonian Institution, PO Box 37012, MRC 119, Washington, DC; ⁸Laboratory for Earthquake Chemistry, University of Tokyo, Hongo, Bunkyo-ku, Tokyo 113-0033, Japan; ⁹Barrios Tech – Jacobs JETS Contract, NASA-JSC; ¹⁰Jacobs JETS Contract, NASA-JSC; ¹¹Aerodyne Industries – Jacobs JETS Contract, NASA-JSC.

Introduction: The Chelyabinsk meteorite fall on February 15, 2013 attracted much more attention worldwide than do most falls [1-3]. A consortium led by JSC received 3 masses of Chelyabinsk (Chel-101, -102, -103) that were collected shortly after the fall and handled with care to minimize contamination. Initial studies were reported in 2013 [4]; we have studied these samples with a wide range of analytical techniques to better understand the mineralogy, petrology, chronology and exposure history of the Chelyabinsk parent body.

Oxidation and weathering: The samples exhibit little to no oxidation: Mössbauer and Raman spectrometry indicate their fresh character. Mass spectrometry reveals a low but clearly detectable level of terrestrial organics indicating that despite the rapid collection and care of handling some terrestrial contamination is present.

Mineralogy and petrology: Mineralogy, petrology, bulk chemistry and magnetic susceptibility measurements all indicate these masses are LL chondrite material [4]. However, detailed studies show that the masses contain three distinct lithologies (Figure 1). A light colored lithology is LL5 material that has experienced shock at levels near S4, based on mineralogy and textures. A second lithology consisted of shock darkened LL5 material in which the darkening is caused by melt veins, and metal-troilite veins distributed along grain boundaries. A third lithology is an impact melt breccia that formed at high temperatures (~1600 °C), and experienced rapid cooling and degassing of S₂ gas. Shock level S4 was experienced by the LL5 lithology (<20 to 30 GPa) but slightly higher pressures (up to 38 GPa) are suggested by high resolution imaging of textures

in impact melt veins (Figure 2), as compared to results of shock experiments [20].

Chronology: Portions of light and dark lithologies from Chel-101, and the impact melt breccias (Chel-102 and Chel-103) were prepared and analyzed for Rb-Sr, Sm-Nd, and Ar-Ar dating. Results yielded ages that cluster at ~264-312 Ma, 716-1014 Ma, and 1112-1464 Ma thus indicating a complex history of impacts and heating events (Figure 3); these ages are consistent with other studies of Chelyabinsk [1, 5-10]. The wide range of ages indicates the Chelyabinsk parent body did not experience post-shock annealing that other ordinary and R chondrites have experienced [11-14]. In addition, the specific ages do not include a 4.2-4.3 Ga impact event identified in other LL chondrites [15].

Exposure history: Finally, noble gases and Sm isotopic compositions were measured on these same aliquots to determine space exposure history. Most LL chondrites have yielded CRE ages of 6 to 50 Ma [16], but Chelyabinsk yields 1 Ma (also [17-19]). This young age, together with the absence of measurable cosmogenic derived Cr, and a barely detectable neutron capture effect in Sm for Chel-101, indicate that Chelyabinsk may have been derived from a recent breakup event on an NEO of LL chondrite composition.

References: [1] Popova, O.P., et al. (2013) *Science* 342, 1069-1073; [2] Borovicka, J. et al. (2013) *Nature* 503, 235-237; [3] Brown, P.G. et al. (2013) *Nature* 503, 238-241; [4] Righter, K. et al. (2013) *Edmonton MetSoc, MaPS Supplement, #5235*; [5] Lindsay, F. et al., (2015), this meeting; [6] Bouvier, A. (2013) *Large Meteorite Impacts and Planetary Evolution V, #3087*; [7] Beard, S.P. et al. (2014) *45th LPSC*,

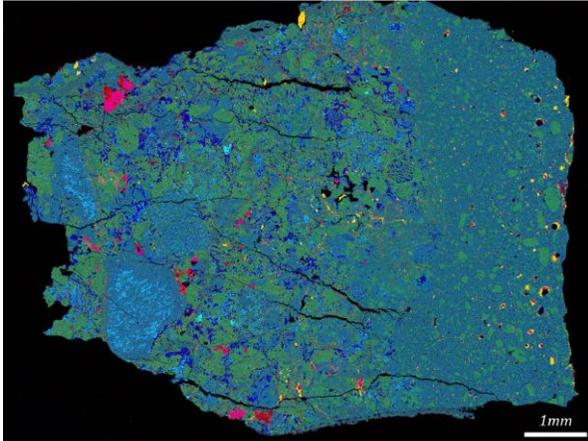


Figure 1: X-ray map of thin section of Chel-102,16. Red = Fe, dark blue / purple = Mg, green = Si, light blue = Ca, magenta = Ni, yellow = S, and white = Ti.

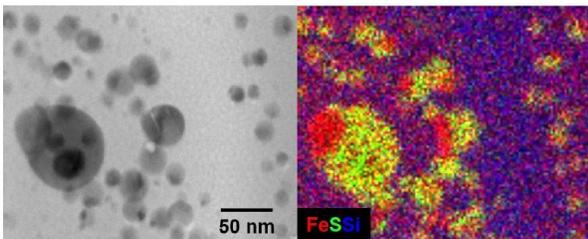
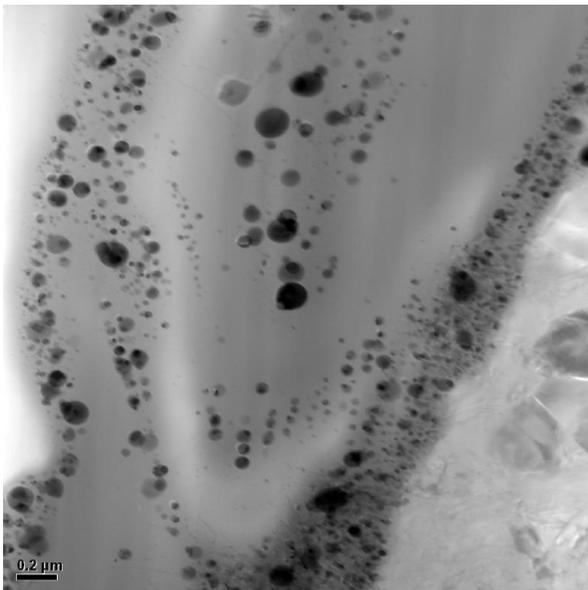


Figure 2 (top): Bright-field STEM image from a FIB section extracted from a region of shock melt in Chelyabinsk. The metal and sulfide inclusions (dark, round grains) are sub-micrometer in size and heterogeneously distributed. Figure 2 (bottom): Bright-field STEM image from a region of shock melt with nanophase inclusions (left) and corresponding composite x-ray map (RGB=Fe S and Si) (right).

#1807; [8] Lapen, T.J. et al., (2014) 45th LPSC, #2561; [9] Galimov, E.M. et al. (2013) *Geokhimiya*, 51, 580–598; [10] Bogolomov, E.S. et al., (2013) *Doklady Earth Sciences* 452, Part 2, 1034–1038; [11] Rubin, A.E. and Jones, R.H. (2003) *MAPS* 38, 1507–1520; [12] Rubin, A.E. (2014) *MAPS* 49, 1057–1075; [13] Righter, K. and Cosca, M. (2014) Goldschmidt Conference; 8-13 Jun. 2014; Sacramento, CA; United States; [14] Rubin, A.E. (2004) *GCA* 68, 673–689; [15] Swindle, T.D. et al. (2014) *Advances in $^{40}\text{Ar}/^{39}\text{Ar}$ Dating: from Archaeology to Planetary Sciences* (eds Jourdan, F., Mark D. F. & Verati, C.), Geological Society, London, Special Publications 378, 333–347; [16] Eugster, O. et al. (2006) *Meteorites and the early solar system II*, 1, 829–851; [17] Haba, M.K. et al. (2014) 45th LPSC, #1732; [18] Nishiizumi, K. et al. (2013) *MAPS Supplement*, #5260; [19] Busemann, H. et al. (2014) 45th LPSC, #2805; [20] Horz, F. et al. (2005) *MAPS* 40, 1329–1346.

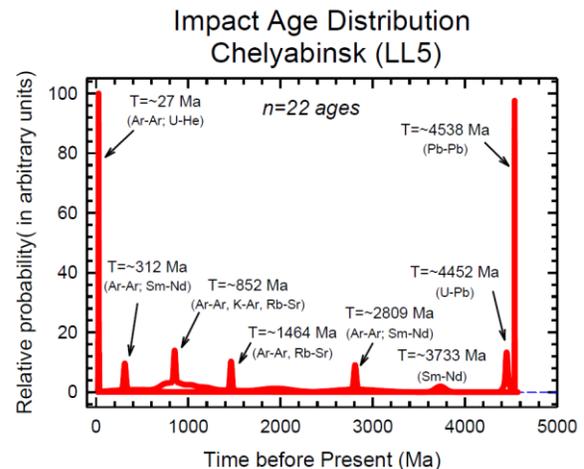


Figure 3: Impact age probability distribution for Chelyabinsk. Red curve represents the summation of the Gaussian distribution of each individual age analysis. The age data of Pb-Pb, U-Pb, Ar-Ar (plateau and “integrated” ages), K-Ar, and U-He [1, 5-8, this study] are more precise and are shown by sharp peaks. The less precise and are shown by broad humps. At least eight impact events (e.g. ~4.53 Ga, ~4.45 Ga, ~3.7 Ga, ~2.8 Ga, ~1.4 Ga, ~312 Ma, and ~27 Ma) are identified for Chelyabinsk. The Sm-Nd and Rb-Sr isotopic systems in Chelyabinsk are very complex and are highly disturbed which is consistent with its having experienced many events of thermal metamorphism and impact resetting after its accretion.