PERCHLORATE ORIGINS IN THE ATACAMA AND ON MARS. D. C. Catling¹, M. W. Claire², K. J. Zahnle³, R. C. Quinn⁴, B. C. Clark⁵, M. H. Hecht⁶, S. P. Kounaves⁷, ¹Earth and Space Sciences/Astrobiology Program, University of Washington, Box 351310, Seattle WA 98195 (<u>dcatling@u.washington.edu</u>). ²Dept. Astronomy/Astrobiology Program, University of Washington. ³NASA Ames Research Center. ⁴SETI Institute/NASA Ames Research Center. ⁵Space Science Institute, Boulder. ⁶JPL/Caltech ⁷Dept. of Chemistry, Tufts University.

Introduction: Perchlorate comprises $\sim 0.3-0.6$ wt% of the soil at the site of NASA's Phoenix Mars Lander [1, 2]. Also, on a molar basis, ClO₄⁻ exceeds Cl⁻ by about an order of magnitude. But what's its origin?

Hypothetical ways to produce Martian perchlorate include photochemistry, electrostatic discharge, and gas-solid reactions [3]. We propose that the Martian environment especially favors perchlorate production associated with atmospheric oxidants. If so, ClO₄ could act as a tracer of past volcanic chlorine and its measurement might be fruitful in returned samples [4] or polar deposits drillcores [5]. However, even on Earth, the exact pathway to produce natural perchlorate is unknown. Consequently, we examine perchlorate formation in the Atacama desert to shed light on Mars because perchlorate levels are comparable to those on Mars only in the Atacama [6]. Isotopic studies indicate that such perchlorate is produced by oxidation involving atmospheric ozone [7-9]. Such pathways terminate in perchloric acid (HClO₄) [10], which can dry deposit.

Pathways to perchlorate: For the terrestrial atmosphere, [10] suggested the pathway:

$$Cl + O_3 + M \rightarrow ClO_3 + M$$
 (1)

$$ClO_3 + OH + M \rightarrow HClO_4 + M$$
 (2)

OH reacts with many species, so (2) relies on enough available ClO₃. OClO can also generate ClO₃ via [11]:

$$OClO + O_3 \rightarrow ClO_3 + O_2 \tag{3}$$

 $OCIO + O + M \rightarrow CIO_3 + M$ (4) where OCIO derives from gas-ice interaction [12] or gas phase chemistry. [13] made another proposal with weakly bound CIO.O₂ and CIO.O₃ intermediates:

$$\text{ClO} + \text{O}_2 + \text{M} \rightarrow \text{ClO}.\text{O}_2 + \text{M}$$
 (5)

$$ClO.O_2 + O_3 \rightarrow ClO.O_3 + O_2 \tag{6}$$

$$\text{CIO.O}_3 + \text{HO}_2 \rightarrow \text{HOCIO}_3 + \text{O}_2 \qquad (7)$$

But $ClO.O_3$ is hypothetical while the stability of $ClO.O_2$ remains questionable [14, 15]. A third model postulates heterogeneous chemistry [16]:

$$\text{ClO} \xrightarrow{\text{H}_2\text{SO}_4} \text{HClO}_4 \text{+HCl} \text{+ products}$$
 (8)

However, measurements have shown that ClO uptake on sulfuric acid is insignificantly slow [17].

Method: We examined an Atacama-like atmosphere with a 1-D photochemical model [18, 19]. A flux of HCl was distributed throughout the troposphere to represent a Cl source from volcanism. Model runs had no rainfall and an Atacama-like tropospheric RH.

Results: Fig. 1 shows deposition fluxes as a function of HCl input. ClO₃ is produced predominantly via (1) and also via (3) and (4); then HClO₄ is produced via (2). Fluxes needed to generate the mass of Atacama salts over the ~10⁷ year age of the deposits are ~10⁸ molecules-NO₃/cm²/s and ~10⁵ molecules-ClO₄/cm²/s [3]. Modeled nitrate fluxes compare favorably (Fig. 1). A match for the required perchlorate flux occurs at HCl fluxes 50-60 times higher than estimates of the globally-averaged HCl flux to the atmosphere, ~0.5 × 10⁹ molecules-HCl/cm²/s [20]. But a much higher local flux is reasonable given the proximity of the Atacama to Andean volcanoes.

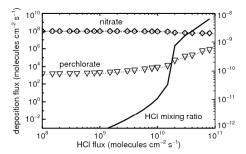


Fig. 1. HCl fluxes (x-axis) result in HCl mixing ratios (solid line) mapped to the R.H. y-axis. Deposition fluxes map to the L.H. y-axis.

Conclusion. We find that the gas phase formation of perchlorate proposed by [10] is feasible for explaining Atacama perchlorate, noting dependence on the uncertain rate of (1). Work is underway with a Mars version of the photochemical model.

References: [1] M.H. Hecht et al. (2009) Science, in press. [2] S.P. Kounaves et al. (2009) JGR, submitted. [3] D.C. Catling et al. (2009) JGR, submitted. [4] MEPAG (2008) Astrobiology, 8, 489-535. [5] M.H. Hecht (2006) 4th Mars Polar Sci Conf, #8096. [6] G.E. Ericksen (1981) USGS. Prof. Paper 1188. [7] H.M. Bao, B.H. Gu (2004) Envir Sci Tech, 38, 5073-5077. [8] J.M. Roberts (2009) Envir Chem, 6, 3-6. [9] N. Kang et al. (2008) Sci Total Envir, 405, 301-309. [10] R. Simonaitis, J. Heicklen (1975) Planet Space Sci, 23, 1567-1569. [11] R.P. Wayne et al. (1995) Atmos Envir, 29, 2677-2881. [12] J.R. McKeachie et al. (2004) J Phys Chem B. 108. 16786-16797. [13] S.S. Prasad, T.J. Lee (1994) JGR, 99, 8225-8230. [14] A. Rauk et al. (1993) J Phys Chem, 97, 7947-7954. [15] D.T. Shindell (1997) J Atmos Chem, 26, 323-335. [16] L.R. Martin et al. (1979) Int J Chem Kin, 11, 543-557. [17] J.P.D. Abbatt (1996) GRL, 23, 1681-1684. [18] K. Zahnle et al. (2006) Geobiology, 4, 271-283. [19] K. Zahnle et al. (2008) JGR, 113, doi:10.1029/2008JE003160. [20] D.M. Pyle, T.A. Mather (2009) Chem. Geol., in press.