

THE STRATA GROUND PENETRATING RADAR: CONSTRAINING THE NEAR SURFACE PROPERTIES OF MARS. J. A. Grant¹, C. J. Leuschen², and P. S. Russell¹, ¹Center for Earth and Planetary Studies, National Air and Space Museum, 6th at Independence SW, Washington, DC, 20560, grantj@si.edu, Dept. of Electrical Engineering and Computer Science, University of Kansas, 1520 West 15th Street, Lawrence, KS 66045.

Introduction: Ground Penetrating Radar (GPR) is a mature technology widely used in terrestrial applications [1-4] and provides an efficient means for non-intrusively defining subsurface radar properties corresponding to structure (e.g., number and size of ejecta blocks, lava tubes, fractures) and stratigraphy to depths of up to tens of meters [e.g., 5-10] (Fig. 1).

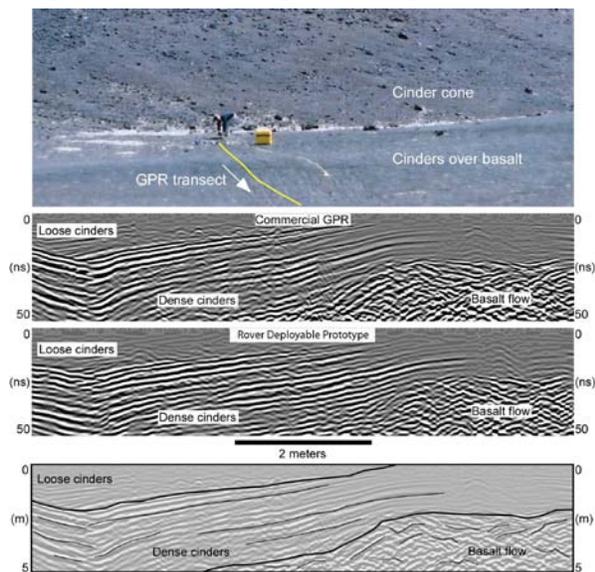


Figure 1. GPR transect across iron-rich volcanic cinders characterized by differing compaction and underlying basalt at Sunset Crater in northern Arizona. Transects completed using a commercial 400 MHz impulse GPR antenna (second from top) and prototype 600 MHz antenna developed for possible future rover deployment on the Moon or Mars (third from top) are shown with interpretation (bottom).

GPR operates by applying a narrow energy pulse through an antenna placed near a surface of interest. The antenna acts as a band-pass filter and emits a single sine wave cycle that is broadcast into the ground at wavelengths ranging from several meters to centimeters (tens of megahertz to a few gigahertz). The longer the incident wavelength, the deeper the expected penetration, but the less ability to resolve closely spaced reflections created by geologic interfaces across changes in physical and electrical properties (e.g., changes in composition). These radar reflections serve as diagnostic “fingerprints” that can help constrain the number and size of buried blocks, thickness and orientation of fractures, and extent of any layers associated with emplacement by different geologic processes (e.g., alluvial vs. volcanic vs. impact). GPR penetration is influ-

enced by composition (e.g., amount of iron or titanium-bearing minerals) and a variety of physical parameters (e.g., grain size and porosity), but the paucity of liquid water, reasonably low loss tangents [e.g. 11], and results from orbital sounding [e.g. 12-13] and Earth-based radars [14] imply GPR should function in most locations on Mars. Although a GPR has not yet flown to Mars, the WISDOM GPR has been selected as part of the PASTEUR payload on the ExoMars rover [15].

The Strata GPR: *Strata* is a low-mass (<3 kg), low-power (<7 W peak), and low-volume (~10,000 cc) impulse GPR (Fig. 2) capable of “peering” beneath the surface of Mars and providing context for the source, setting, and distribution of detections made by other instruments [16-17]. *Strata* design and development has been funded by multiple NASA programs and the instrument has been field tested and is ready for inclusion in the payload on a future mission to Mars.

Strata (Fig. 2) consists of a low mass and low power digital processor unit and a set of loaded dipole antennas operating at 400 MHz (75-cm wavelength). With its high dynamic range (~110 dB), *Strata* should be capable of probing 10-15 m into the subsurface of Mars with a nominal resolution of ~30 cm capable of defining stratigraphy and structure required to characterize ancient habitable setting on Mars.

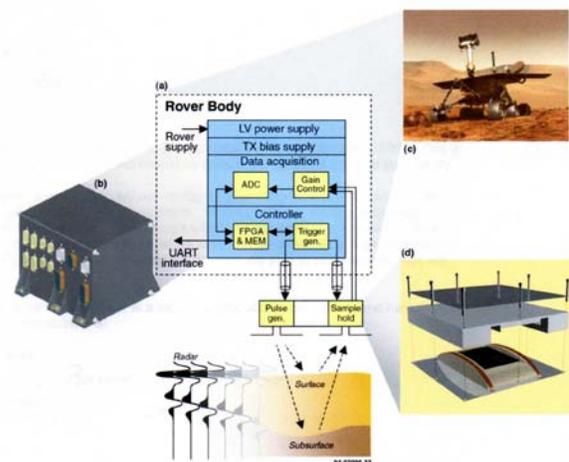


Figure 2. Basic components and deployment scheme for the *Strata* GPR. (a) *STRATA* block diagram with (b) a typical APL instrument digital processor unit (c) a view of the rover, and (d) packaging concept for the antenna assembly that would be mounted on the underside of the rover. DPU = digital processor unit.

Strata's antenna can be mounted above the surface on the belly of the rover (protrudes <10 cm) and operates well in high-loss settings characterized by iron-rich substrates (Fig. 1). Positioning the antenna beneath the rover and above the ground ensures it is not a hazard to rover operations, and it can easily be placed to avoid obstructing the field of view of other instruments. Data requirements are low, with ~0.6 MB/day expected (for 100 meter traverses).

Science With *Strata*: *Strata* can constrain the nature and variability of stratigraphy and surficial deposits along rover traverses and relate them to geologic setting. *Strata* can also help identify potential hazards to rover trafficability and assist in future human exploration of Mars via definition of the extent/thickness of near-surface fines and helping to pinpoint ice deposits.

Strata will define shallow stratigraphy, thereby exposing "virtual" outcrops in the near-surface that can be compared to more scattered, local outcrops and used in evaluating geologic setting. Without knowledge of how data from other rover instruments relate to the overall geologic evolution of a site, samples obtained for analysis can lack the context needed for accurate interpretation. Hence, sub-surface *Strata* data are essential in targeting samples for other remote and in situ investigations and can help locate the best samples for sample return to Earth.

Much of Mars is mantled by surficial deposits that modify, mask, and impede characterization of underlying deposits and structures diagnostic of high-priority target environments and can cloud interpretation of the origin of intriguing geologic settings. For example, much of the nominal mission of both MER rovers was spent trying to get information about the nature and origin of any bedrock at the landing sites. This required long traverses to locate, characterize, and access in situ rocks that might record a water-lain history. Had *Strata* been part of the MER payload, it would have made defining the distribution and character of bedrock more straightforward and helped guide the rovers quickly to samples best suited to meeting mission objectives.

Ongoing work with GPR in impact, arid alluvial, eolian, volcanic, and polar settings all contribute to a growing database that can be used to distinguish settings that may be encountered on Mars during future exploration. For example, the interpreted distribution of blocks in impact ejecta at Meteor Crater, AZ, using a 400 MHz antenna (the same λ of 75 cm as *Strata*) is 1.5-3.0 blocks per m^3 in the uppermost 1 m of the subsurface (and 0.5-1.0 blocks per m^3 in the uppermost two meters), which is close to the *in situ* measured block distribution of 2.0-3.0 blocks larger than 0.25-0.30 m per m^3 [18]. This is roughly the detection limit to be expected from the $\lambda/3$ resolution approximation

of the radar wavelength and indicates that the 400 MHz GPR is characterizing the block population in ejecta. This population can then be compared to the distribution of blocks and radar stratigraphy mapped in alternate geologic settings or at other impact craters. Collectively, results from terrestrial analogs for Martian settings will contribute to understanding the likely range in radar properties of the upper 10-15 m of the Mars subsurface. Results suggest that interpretation of geologic setting using characteristic radar signatures and quantitative constraint of subsurface properties is possible using *Strata*.

References: [1] GPR 1994 (1994) Proc. 5th Int'l Conf. on GPR, 1294 pp., Waterloo Centre for Groundwater Research, Kitchener, Ontario, Canada. [2] GPR 1998 (1998) Proc. 7th Int'l Conf. on GPR, 786 pp., Radar Systems and Remote Sensing Laboratory, Lawrence, KS. [3] GPR 2000 (2000) Proc. 8th Int'l Conf. on GPR, D.A. Noon, G.F. Stickley and D. Longstaff, Eds., SPIE 4084, 908 pp., Gold Coast, Australia. [4] GPR 2002 (2002) Proc. 9th Int'l Conf. on GPR, S. Koppenjan and H. Lee, Eds., SPIE 4758, 734 pp., Santa Barbara, CA. [5] Ulriksen, C. P. F. (1982), Application of impulse radar to civil engineering: Ph.D. Thesis, University of Technology, Lund, Sweden, 175p. [6] Pilon, J. A., et al. (1991), JGR, 96, 15,563-15,576. [7] Grant, J. A., and Schultz, P. H. (1993), JGR, 98, 15,033-15,047. [8] Grant, J. A., and Schultz, P. H. (1994), Erosion of ejecta at Meteor Crater: Constraints from ground penetrating radar, in Proc. 5th Int'l Conf. on GPR, Univ. of Waterloo, Ontario, Canada, pp. 789-803. [9] Grant, J. A. et al. (2003), JGR., 108, 10.1029/2002JE001856. [10] Grant, J. A. et al. (2004), JGR, 109, 10.1029/2003JE002232. [11] Campbell, B.A., et al. (2008), JGR, 113, doi:10.1029/2008JE003177. [12] Picardi, G., et al. (2005) Science 310, 1925-1928, doi: 10.1126/science.1122165. [13] Seu, R., et al. (2007) JGR, 112, doi: 10.1029/2006JE002475. [14] Harmon, J. K. and Nolan, M. C. (2007), 7th Int. Conf. Mars, abs. 3136, Pasadena, CA. [15] Ciarletti, V. (2006) *EGU Geophys. Res. Abs.*, 8, 09136. [16] Leuschen, C. J., et al. (2005), Eos Trans. AGU 86, P31C-0210. [17] Williams, K.K., et al. (2005), Eos Trans. AGU 86, P31C-0209. [18] Russell, P. S. et al. (2012), LPSC 43, abs. 1612, LPI, Houston, TX.