REDOX CHEMISTRY AND THE ORIGIN OF ACIDITY ON THE ANCIENT SURFACE OF MARS. J. A. Hurowitz^{1*}, W. W. Fischer², N. J. Tosca³, and R. E. Milliken¹, ¹Jet Propulsion Laboratory, California Institute of Technology, 4800 Oak Grove Drive, Pasadena, CA, 91109, ²Division of Geological and Planetary Sciences, 107 North Mudd Laboratory, California Institute of Technology, Pasadena, CA 91125, ³Department of Earth Sciences, University of Cambridge, Cambridge CB2 3EQ, United Kingdom, *joel.a.hurowitz@jpl.nasa.gov.

Introduction: *In-situ* and orbital exploration have demonstrated that surface waters at Meridiani Planum, Mars were acidic [1-4]. However, the origin of this acidity is unknown. Constrained by chemical and mineralogical analyses from the Mars Exploration Rover *Opportunity* [5-8], we show that Fe-oxidation and Fe³⁺-mineral precipitation yields an excess of acid relative to the amount of titrant available in outcrop [9]. Our results indicate that Fe²⁺-bearing subsurface waters, buffered to circum-neutral pH and anoxia, were subject to rapid oxidation and acidification upon exposure to O_2 and/or UV light (**Fig. 1**).

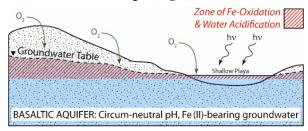


Fig. 1: Schematic cross-section showing the oxidation and acidification of groundwater at Meridiani Planum.

Approach: In-situ measurements of Meridiani outcrop chemistry and mineralogy make it possible to quantify the amount of acid consumed or produced during Fe-oxidation and Fe³⁺ mineral precipitation. The acid produced can then be "titrated" against the available base anion content measured in outcrop (i.e., SO_4^{2-} , Cl^- , PO_4^{3-} , and CO_3^{2-}), yielding a net proton balance reflective of the parent fluid. For our titration method, we sum the number of moles of H⁺ generated during precipitation of the measured quantity and distribution of jarosite, hematite, and schwertmannite, using data collected through sol 548 of the Opportunity mission on 19 Meridiani Planum outcrop targets. Our approach is based on a straightforward accounting of the number of moles of acid produced when Fe3+ minerals are formed from dissolved Fe²⁺, e.g.,:

(1)
$$Fe^{2+}$$
 (aq) + 0.25O₂ (aq) + H⁺ = Fe^{3+} (aq) + 0.5H₂O
(2) Fe^{3+} (aq) + 0.667SO₄²⁻ + 0.33K⁺ + 2H₂O = 0.33KFe₃(SO₄)₂(OH)₆ (jarosite) + **2H**⁺

Results: For all outcrop analyses there is an excess of H⁺ generated in forming the observed secondary Fe³⁺ mineral phases relative to the available titrant in

outcrop (**Fig. 2**). Our calculations imply that so long as redox conditions in the aquifer were conducive to the transport of Fe²⁺(aq), oxidation and formation of jarosite, hematite, and schwertmannite would have resulted in the generation of low-pH fluids at the site of Fe³⁺-mineral precipitation (**Fig. 1**). Accordingly, input of additional acid volatiles (e.g., SO₂, H₂SO₄) at the site of sediment formation [10, 11] is not required and SO₄²⁻ can be considered a background constituent of the aquifer fluid.

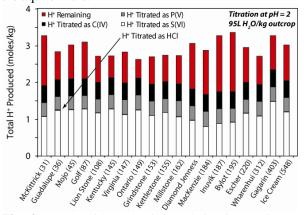


Fig. 2: Total acid production and titration against a pH=2 base species distribution for analyses collected between sols 31 and 548 on abraded outcrop targets.

Finally, because gaseous H_2 is a by-product, Feoxidation processes have an impact on the redox state of the atmosphere. We will show that the Martian sedimentary record provides quantitative insight into the magnitude and timing of atmospheric H_2 loss processes required for oxidation at Meridiani Planum.

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