**WIND-ERODED FLOOR DEPOSITS IN NOACHIAN DEGRADED CRATERS ON MARS.** R. P. Irwin III<sup>1,2</sup>, F. Pendrill<sup>3</sup>, T. A. Maxwell<sup>2</sup>, A. D. Howard<sup>4</sup>, and R. A. Craddock<sup>2</sup>, <sup>1</sup>Planetary Science Institute, 1700 E. Ft. Lowell Rd., Suite 106, Tucson, AZ 85719, irwin@psi.edu. <sup>2</sup>Center for Earth and Planetary Studies, National Air and Space Museum, Smithsonian Institution, MRC 315, 6th St. at Independence Ave. SW, Washington DC 20013-7012. <sup>3</sup>University of Gothenburg, SE-405 30, Gothenburg, Sweden. <sup>4</sup>Department of Environmental Sciences, University of Virginia, Charlottesville VA 22904.

Introduction: The role of water in Noachian crater degradation is key to understanding the paleoclimate of early Mars. Most global- to regional-scale geologic maps suggested widespread volcanic resurfacing of impact crater floors in the early Hesperian [1,2]. Large degraded craters often have flat floors that are consistent with volcanic resurfacing, as the Mars Exploration Rover Spirit found in Gusev crater [3]. Many Noachian craters regardless of size have floors with high thermal inertia [4]. However, the topography of degraded crater rims and interior deposits indicates a dominant role for fluvial erosion and deposition, particularly in craters <60 km in diameter [5,6].

We examined >1000 floors of Noachian craters from 0–30°S, 0–165°E to identify friable deposits that (unlike basalts) were susceptible to aeolian deflation. We used visible-wavelength orbital imaging at 0.25–18 m/pixel from the High Resolution Imaging Science Experiment, Mars Orbiter Camera, Context Camera, and Thermal Emission Imaging System.

Erosional Resistance of Crater Floors: The morphology of Noachian crater floors varies greatly, suggesting diverse fill materials and/or weathering and erosion processes [7]. Some floor materials retain abundant small, post-depositional craters, but they are otherwise smooth. At the opposite extreme are knobby floors (Fig. 1), some of which have exposed layers. In many cases we noted light-toned, seemingly friable material overlain by a more resistant, darker unit. During post-depositional modification, deflation of the darker-toned surface layer removes much of the smallcrater population and may expose the underlying lighttoned layer, whereas less effective deflation leaves patchy exposures of deeper materials. These friable, wind-eroded deposits are confined to the floors of craters and intercrater basins.

**Spatial Distribution of Wind-eroded Crater Floors:** Knobby, deflated crater floors are concentrated in a southern band within the study area, except between 60–90°E, where we found no north-to-south increase in floor erosion. Light-toned outcrops on basin floors are concentrated south of 20°S. The area from 20–25°S, 25–100°E contains about a third of the outcrops of light-toned material in the study area, with concentrations near the rim of Hellas and southwest of Huygens crater (~20–25°S, 40–50°E). The observed

concentration of fracturing in light-toned outcrops in these areas may be due to better image coverage.

**Discussion:** Possible reasons for the diverse Noachian crater-floor morphology are: 1) different resurfacing processes around the Noachian/Hesperian transition; 2) different intensity of post-Noachian aeolian modification; and/or 3) different surface weathering rates, possibly related to material properties. The concentration of knobby terrain and light-toned layers in crater or intercrater basins suggests that these units are not an airfall mantle, and that any enhanced weathering processes must have been focused in the basins. Winderoded crater floors appear concentrated in low-albedo, dust-free regions, so increased aeolian efficiency may be partly responsible.

**Conclusions:** Many Noachian craters have friable, wind-eroded deposits restricted to their floors, some with deeper light-toned material exposed, suggesting aqueous depositional environments. These craters are concentrated in low-albedo, dust-free highland regions. In many cases, basaltic volcanism was not the last significant resurfacing process in these craters.

**References:** [1] Scott D. H. and Carr M. H. (1978) USGS Map I-1083. [2] Scott D. H. et al. (1986–1987) USGS Map I-1802. [3] Arvidson R. A. et al. (2006) *JGR*, 111, E02S01, doi:10.1029/2005JE002499. [4] Putzig N. E. and Mellon M. T. (2007) Icarus, 191, 68–94. [5] Craddock R. A. et al. (1997) *JGR*, 102, 13,321–13,340. [6] Craddock R. A. and Howard A. D. (2002) *JGR*, 107(E11), 5111, DOI:10.1029/2001JE001505. [7] Irwin R. P. et al. (2009) *LPSC* 40, 2358.

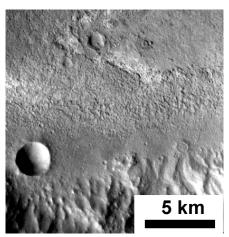


Fig. 1. Gullied Noachian crater wall and etched floor, THEMIS V07923006, 28.1°S, 129.8°E.