

**<sup>26</sup>Al-<sup>26</sup>Mg SYSTEMATICS OF ALLENDE AOAs.**

M.B. Olsen<sup>1</sup>, K. Larsen<sup>1</sup>, C. Paton<sup>1</sup>, D. Wielandt<sup>1</sup>, A.N. Krot<sup>1,2</sup> and M. Bizzarro<sup>1</sup>. <sup>1</sup>Centre for Star and Planet Formation, Natural History Museum of Denmark, University of Copenhagen, Copenhagen 1350 Denmark. <sup>2</sup>HIGP, University of Hawai'i at Manoa, HI, USA.

Amoeboid Olivine Aggregates (AOAs) are irregularly shaped mm to cm sized refractory objects present in most carbonaceous chondrite groups. AOAs are believed to have formed in an <sup>16</sup>O-rich gaseous reservoir as aggregates of high temperature nebular condensates consisting of forsterite, Fe-Ni metal and a refractory Ca-Al-rich component, and have escaped extensive melting [1]. Recent work on the Ti isotope composition of early solar system solids, asteroids and planets [2] indicates that AOAs and CAIs have identical <sup>50</sup>Ti excesses, suggesting a genetic link between these two types of refractory materials. Thus, similarly to CAIs, AOAs can provide information regarding the earliest evolution of the solar protoplanetary disk. Moreover, available <sup>26</sup>Al-<sup>26</sup>Mg data for AOAs obtained by in situ methods (SIMS) suggests that these solids formed with an initial <sup>26</sup>Al/<sup>27</sup>Al value of  $\sim 5 \times 10^{-5}$  [3], consistent with the canonical abundance of <sup>26</sup>Al in the early solar system.

Here, we report the preliminary results of a study aimed at characterizing the petrology and <sup>26</sup>Al-<sup>26</sup>Mg systematics of a number of AOAs from the Allende CV chondrite, in order to better understand the duration of the AOA-forming event and its relationship to the CAI-forming event. AOAs were selected from polished slabs of the Allende meteorite and, following petrographic characterization, sampled by means of micro-drilling. Following Mg purification by ion exchange chromatography, the Mg isotope composition and Al/Mg ratios were determined by high-resolution multiple collection plasma source mass spectrometry. Novel analytical protocols allow for measurements of  $\mu^{26}\text{Mg}^*$  (per  $10^6$  deviation from terrestrial <sup>26</sup>Mg-/<sup>24</sup>Mg) and Al/Mg ratios in silicate materials to higher precision than previously possible, namely with an external reproducibility and accuracy of 2.5 ppm and 0.2%, respectively.

Four AOAs with <sup>27</sup>Al/<sup>24</sup>Mg ratios ranging from 0.115 to 0.874 define a line in  $\mu^{26}\text{Mg}^*$  vs. <sup>27</sup>Al/<sup>24</sup>Mg space with a slope corresponding to an initial <sup>26</sup>Al/<sup>27</sup>Al of  $(5.0 \pm 0.3) \times 10^{-5}$ , and an intercept of  $-6.5 \pm 8.0$  ppm. The uncertainty of the slope defined by the Allende AOAs sampled so far, corresponds to an age uncertainty of  $\sim 60,000$  years, thereby supporting a short duration for the AOA-forming event. In contrast to recent high-precision <sup>26</sup>Al-<sup>26</sup>Mg data for Efremovka AOAs and CAIs [4], the Allende AOAs do not define a statistically significant isochron (MSWD=4.5), indicating a more severe disturbance of the <sup>26</sup>Al-<sup>26</sup>Mg systematics in Allende inclusions as compared to Efremovka. However, similarly to Efremovka AOAs, the regression of the Allende AOAs intercepts the solar <sup>27</sup>Al/<sup>24</sup>Mg ratio of 0.1 at a  $\mu^{26}\text{Mg}^*$  value of  $29 \pm 8$  ppm, that is, significantly higher than the terrestrial  $\mu^{26}\text{Mg}^*$ . This observation may be interpreted as reflecting a higher abundance of <sup>26</sup>Al in the CAI-AOA forming reservoir compared to Earth's precursor material, implying a heterogeneous distribution of <sup>26</sup>Al in the solar protoplanetary disk and, a more complicated significance of <sup>26</sup>Al-<sup>26</sup>Mg data, as an extra variable needs to be constrained [4].

**References:** [1] Komatsu, M. *et al.* 2001. MAPS 36: 629-641. [2] Trinquier, A. *et al.* 2009. Science 64:185-239. [3] Sugiura, N. and Krot, A.N. 2007. MAPS 42: 1183-1195. [4] Larsen, K. *et al.* 2010. GCA, this meeting.