

NOBLE GASES IN COSMIC SPHERULES AND MICROMETEORITES FROM THE TRANSANTARCTIC MOUNTAINS.

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Introduction: At MPI Mainz we have initiated a comprehensive survey of noble gases in micrometeorites. While He and Ne in micrometeorites are generally dominated by solar contributions, the inventory of heavy primordial noble gases is not as well characterized. We hope to fill this gap with particular focus on the diagnostic isotopic composition of Xe. While our study will eventually include material from a variety of collections (e.g. the CONCORDIA collection [1]), our first results are from “large” PNRA (Programma Nazionale delle Ricerche in Antartide, Italy) micrometeorites collected from the Victoria Land, Transantarctic Mountain micrometeorite traps [2].

Samples and Analytics: 51 micrometeorites (MM's) - 26 cosmic spherules (CS), 11 scoriaceous (Sc), 14 unmelted (Un) - in the 600-800 μm size range were extracted from traps #45 b and c on the top of Miller Butte, and split into aliquots for combined petrographic and noble gas analyses. For the latter, aliquots were heated using a CO₂ laser. The analyses were performed using our multicollector Nu Instruments noble gas mass spectrometer [3].

Preliminary Results: At present, we have obtained noble gas results for three CS and one Sc micrometeorites. Our Ne and Ar results (corrected for isobaric interferences, but not extraction blank) are listed in Table 1.

Table 1. Ne and Ar results for MM's. Sizes are given in μm . Weights are given in μg , noble gas amounts - not corrected for blank - in 10^{-12} cc units. Uncertainties in the last digits.

Sample	Size	Weight	²² Ne	²⁰ Ne/ ²² Ne	²¹ Ne/ ²² Ne	³⁶ Ar	⁴⁰ Ar/ ³⁶ Ar
CS 45c21	720 x 556	337.9	0.41	9.11 (37)	0.029 (2)	0.40	319 (17)
CS 45c17	630 x 476	126.3	0.16	9.34 (1.01)	0.038 (6)	2.36	280 (9)
CS 45c16	465 x 324	105.1	0.21	8.85 (85)	0.027 (4)	0.98	292 (10)
Sc 45b22	268 x 191	9.2	0.49	10.46 (45)	0.062 (4)	1.04	128 (5)

As found in previous studies (e.g. [4, 5]) noble gas abundances in CS are mostly low and essentially have a terrestrial air signature (e.g. ²⁰Ne/²²Ne ~9.8, ²¹Ne/²²Ne ~0.029). In contrast, the Sc micrometeorite shows ²⁰Ne/²²Ne and (even clearer) ²¹Ne/²²Ne ratios higher than air, indicative of the presence of solar and spallogenic Ne. Using the production rates for IDPs from table 3 in [6], the spallogenic ²¹Ne of $\sim 1.6 \times 10^{-9}$ cc/g would be produced at 1 AU by SCR primaries in ~ 0.17 Ma, for production by GCR primaries, the inferred exposure would be ~ 4 Ma. The concentration of trapped (solar) ²⁰Ne of $\sim 5 \times 10^{-7}$ cc/g is about an order of magnitude lower than the geometric mean for Un micrometeorites from Yamato Mountains and Dome Fuji, but distinctly higher than in CS (see table 4 in [4]). The low ⁴⁰Ar/³⁶Ar ratio for the Sc micrometeorite (~ 128) indicates the presence of trapped ³⁶Ar, at an abundance of $\sim 6 \times 10^{-8}$ cc/g. This is most likely “planetary” type Ar, carried primarily by phase Q [7].

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References: [1] Duprat J. et al. (2007) *Advances in Space Research* 39:605-611. [2] Rochette P. et al. (2008) *Proceedings of the National Academy of Sciences* 105:18206-18211. [3] Ott U. et al. (2010) *Meteoritics & Planetary Science* 45:A158. [4] Osawa T. et al. (2010) *Meteoritics & Planetary Science* 45:1320-1339 [5] Marty B. et al. (2005) *Meteoritics & Planetary Science* 40:881-894. [6] Kehm K. et al. (2006) *Meteoritics & Planetary Science* 41:1199-1217. [7] Lewis R. S. et al. (1975) *Science* 190:1251-1262.