

COSMOGENIC RECORDS IN 18 ORDINARY CHONDRITES FROM THE DAR AL GANI REGION, LIBYA: I NOBLE GASES. L. Schultz¹, L. Franke¹, K. C. Welten², K. Nishiizumi² and A. J. T. Jull³, ¹Max-Planck-Institut für Chemie, Postfach 3060, D-55020 Mainz, Germany (schultz@mpch-mainz.mpg.de), ²Space Sciences Laboratory, University of California, Berkeley, CA 94720-7450, USA (kcwelten@uclink4.berkeley.edu), ³NSF Arizona AMS Facility, University of Arizona, Tucson, AZ 85721, USA.

Introduction: In the last decade thousands of meteorites have been recovered from hot deserts in the Sahara and Oman. One of the main meteorite concentration surfaces in the Sahara is the Dar al Gani plateau in Libya, which covers a total area of ~8000 km². More than 1000 meteorites have been reported from this area. The geological setting, meteorite pairings and the meteorite density of the Dar al Gani (DaG) field are described in more detail in [1].

In this work we report concentrations of the noble gas isotopes of He, Ne, Ar as well as ⁸⁴Kr and ¹³²Xe in 18 DaG meteorites. In a separate paper we will report the cosmogenic radionuclides [2]. We discuss the thermal history and cosmic-ray exposure (CRE) history of these meteorites, and evaluate the effects of the hot desert environment on the noble gas record.

Samples: We selected 13 H-chondrites, 4 L-chondrites and 1 LL-chondrite for noble gas and radionuclide analyses. Only one of these meteorites belongs to a shower, whereas two others belong to two different meteorite pairs [1]. The weathering grade ranges from W1 (DaG 062/908) to W5 for DaG 388, the shock grade from S1 (DaG 907) to S4 (DaG 908).

Experimental procedures and results: Concentrations and isotopic ratios of He, Ne and Ar were measured in chips of ~100 mg at the MPI, as described in [3]. The measured noble gas concentrations, given in Table 1, are a mixture of trapped components, a radiogenic component and a cosmogenic component, which will be discussed separately.

Trapped noble gases. Most samples contain significant amounts of trapped Ne, Ar, Kr and Xe, which generally increase from weathering grade 1 to 4, most notably for ⁸⁴Kr and ¹³²Xe. The ratio of trapped ¹³²Xe/⁸⁴Kr ranges from 0.1-0.4, which is much closer to the value of 0.073 in water (at 0 °C) than to the average planetary value of ~1.7 [4]. These two observations and the lack of ⁴He concentrations >2 x 10⁻⁵ cm³ STP/g indicate that the majority of the trapped noble gases is atmospheric. We therefore assume a ²⁰Ne/²²Ne ratio of 9.8 and a ³⁶Ar/³⁸Ar ratio of 5.32 for the trapped Ne and Ar components [3].

Radiogenic noble gases. The concentrations of radiogenic ⁴He are deduced from the measured ⁴He concentration by subtracting the cosmogenic component, assuming (⁴He/³He)_c=6 [5]. The concentrations of radiogenic ⁴⁰Ar are deduced from the measured ⁴⁰Ar concentration by subtracting a trapped atmospheric component, assuming (⁴⁰Ar/³⁶Ar)_{tr} = 250±50 [6]. Radiogenic ⁴He ranges from 0.03–1.80 x 10⁻⁵ cm³ STP/g, whereas radiogenic ⁴⁰Ar ranges from 0.2–6.0 x 10⁻⁵ cm³ STP/g. The radiogenic ⁴He and ⁴⁰Ar concentra-

tions seem to correlate with shock stage, with consistently high values for samples with S1-S2 and lower values for samples with S3-S4, consistent with previous observations [7]. Surprisingly, none of the five samples with low radiogenic ⁴⁰Ar (below 4.0 x 10⁻⁵ cm³ STP/g), are L-chondrites.

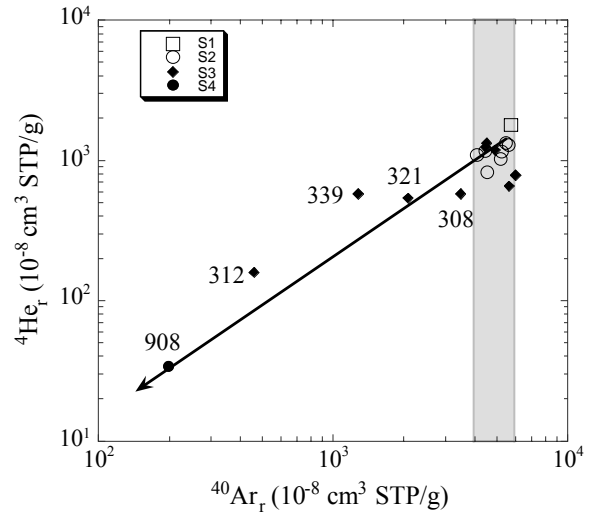


Fig. 1. Correlation of radiogenic ⁴He and ⁴⁰Ar with shock stage. Most DaG samples with S3 and S4 show significant losses of radiogenic gases.

Cosmogenic noble gases. We determined the concentrations of cosmogenic ³He, ²¹Ne, ²²Ne and ³⁸Ar as in [3]. In many DaG samples, the concentration of cosmogenic ³⁸Ar is difficult to determine due to large amounts of trapped Ar: in a few cases the cosmogenic ³⁸Ar component is less than 10% of the total ³⁸Ar. The concentration of cosmogenic ²¹Ne is much more reliable, because the amount of trapped ²¹Ne is usually negligible (<1%). However, for samples with large amounts of trapped Ne, the cosmogenic ²²Ne/²¹Ne ratio, which is used as a shielding parameter, becomes less reliable. Three DaG samples show highly elevated ²⁰Ne/²²Ne ratios of 3.5-4.5 relative to a typical ²⁰Ne/²²Ne ratio of ~0.84 for cosmogenic Ne. For these samples we adopted a ²²Ne/²¹Ne ratio of 1.11, corresponding to average shielding (Table 1).

Exposure ages. We calculated the CRE ages from the cosmogenic ³He, ²¹Ne and ³⁸Ar concentrations and production rate methods which use the ²²Ne/²¹Ne ratio as a shielding parameter [8]. The ²¹Ne ages show typical values for ordinary chondrite in the range from ~1 Myr to over 40 Myr. Surprisingly, only one of the 13 DaG H-chondrites shows a CRE age coinciding with the main cluster at 7 Myr, whereas 8 meteorites show

exposure ages ≤ 5 Myr and two coincide with the peak at ~ 30 Myr. This is partly due to the selection of samples, which was done using cosmogenic gases to avoid pairing. It is also likely that some of the low CRE ages were underestimated, since many samples show $^{22}\text{Ne}/^{21}\text{Ne}$ ratios < 1.10 . At these high shielding conditions, the simple relationship between the ^{21}Ne production rate and the $^{22}\text{Ne}/^{21}\text{Ne}$ ratio is not valid, resulting in an overestimation of the production rate for very large objects. In order to obtain reliable ages we will calculate $^{10}\text{Be}/^{21}\text{Ne}$ ages for these samples [2].

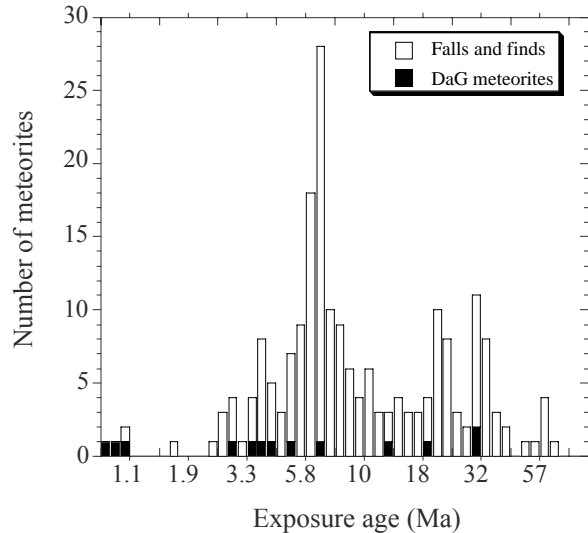


Fig. 2. CRE ages of 13 H-chondrites from the DaG region compared to those of 190 falls and finds [9].

The ^3He ages show generally good agreement (within 20%) with the ^{21}Ne ages, even for meteorites of weathering grade W4 and W5. This implies that moderate weathering in hot deserts does not lead to significant loss of cosmogenic ^3He . Nevertheless, six samples show ^3He ages significantly lower than the corresponding ^{21}Ne ages. For one or two samples with large

amounts of trapped Ne, this is mainly due to the uncertainty in the $^{22}\text{Ne}/^{21}\text{Ne}$ ratio, which strongly affects the ^{21}Ne production rate and thus the ^{21}Ne age. However, for four other samples (DaG 308, 321, 339 and 908), the low ^3He ages correlate with loss of radiogenic gases. This correlation suggests that cosmogenic ^3He and radiogenic ^4He and ^{40}Ar were lost either by shock heating during impact on the parent body or by solar heating in space. The first scenario requires a previous CRE on the parent body. Results of cosmogenic radionuclide measurements indicate that DaG 908 had such a complex exposure history [2], whereas for the other samples with low cosmogenic ^3He , losses are most likely due to solar heating.

Conclusions: The 18 meteorites studied represent 18 individual falls. The trapped noble gases are dominated by atmospheric gases, which increase with the degree of weathering, especially for Kr and Xe. Five of the samples show significant losses of radiogenic ^4He and ^{40}Ar , which correlate with shock effects rather than with terrestrial weathering. Eight of the 13 H-chondrites show CRE ages below 5 Myr, whereas only one coincides with the peak at ~ 7 Myr. The low CRE ages may be an artifact of the high shielding in many samples as judged from the low $^{22}\text{Ne}/^{21}\text{Ne}$ ratios. Loss of cosmogenic ^3He in six samples is not due to weathering but to solar heating or impact-related heating.

References: [1] Schlüter J. et al. (2002) MAPS 37, 1079. [2] Welten K. et al. (2003) LPSC 34, this volume. [3] Scherer P. et al. (1998) MAPS 33, 259 [4] Scherer P. et al. (1994) In: *Noble Gas Geochem. and Cosmochem.*, pp. 43-53. [5] Alexeev V. (1998) MAPS 33, 145. [6] Welten K. et al., MAPS 38 (in press). [7] Schultz L. and Stöffler D. (1993) MAPS 28, 432. [8] Eugster O. (1988) GCA 52, 1649. [9] Graf T. and Marti K. (1995) JGR 100, 21247.

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Table 1. Noble gas concentrations (10^{-8} cm³ STP/g) and ^3He , ^{21}Ne and ^{38}Ar exposure ages (in Myr) of 18 DaG meteorites.

DaG	Class	^3He	^4He	^{20}Ne	^{21}Ne	^{22}Ne	^{36}Ar	^{38}Ar	^{40}Ar	$^3\text{He}/^{21}\text{Ne}$	$^{22}\text{Ne}/^{21}\text{Ne}$	T3	T21	T38
304	H6,S3,W4	1.07	789	2.34	0.33	0.53	16.47	3.16	9577	3.20	1.11	0.7	1.1	1.5
308	H6,S3,W2	2.92	595	2.31	1.17	1.35	6.50	1.36	5093	2.49	1.041	1.8	2.8	3.2
311	H6,S3,W3	6.92	1220	2.61	1.63	1.76	3.37	0.79	6174	4.26	1.044	4.2	3.7	3.7
312	H6,S3,W2	40.19	400	8.83	9.35	10.49	1.98	1.31	774	4.30	1.121	25.0	31.8	25.2
321	H5,S3,W3	5.73	571	4.32	2.29	2.72	1.49	0.49	2420	2.50	1.091	3.5	6.8	5.3
322	H4,S2,W2	30.42	1332	4.75	4.05	5.01	5.78	1.67	6548	7.51	1.221	19.4	19.6	18.7
336	H5/6,S2,W4	5.33	1317	1.43	0.95	1.13	6.51	1.33	7196	5.62	1.138	3.3	3.5	2.9
339	H5,S3,W3	3.29	596	8.83	1.57	2.45	5.64	1.23	2656	2.10	1.11	2.0	5.0	4.5
343	H4,S2,W4	6.51	1358	1.07	0.81	1.07	4.41	0.93	6700	8.00	1.288	4.2	4.4	3.9
388	H5/6,S2,W5	44.6	1429	10.51	11.41	12.31	2.74	1.55	4929	3.89	1.077	27.3	31.5	26.0
904	H6,S2,W3	1.14	834	0.36	0.39	0.41	1.17	0.24	4815	2.95	1.075	0.7	1.0	0.5
907	H6,S1,W3	22.66	1919	4.27	4.61	5.03	1.75	0.90	6051	4.92	1.090	14.0	13.6	14.6
908	H6,S4,W1	0.22	36	2.05	0.27	0.48	5.26	1.00	1512	0.80	1.11	0.1	0.9	0.8
330	L5,S2,W3	12.52	1098	2.47	2.52	2.91	12.73	2.29	8380	4.96	1.154	7.9	9.1	-
341	L6,S3,W3	17.71	761	3.07	3.33	3.85	2.45	0.78	6169	5.33	1.157	11.1	12.1	9.8
342	L5-6,S2,W3	58.20	1437	15.69	17.10	18.31	3.80	2.06	4784	3.40	1.069	35.7	41.9	36.1
906	L6,S3,W2	8.30	1374	1.85	0.87	1.22	6.25	1.25	6048	9.54	1.296	5.4	4.8	3.0
062	LL5-6,S3,W1	19.58	1350	3.40	3.60	4.09	2.05	0.86	6007	5.45	1.134	12.2	12.0	14.0

Low ^3He ages which are due to loss of cosmogenic ^3He are printed in italics.