

**$^{39}\text{Ar}$ - $^{40}\text{Ar}$  Dating of Martian Shergottite, DaG 476. Jisun Park<sup>1,2</sup>, Donald D. Bogard<sup>1</sup> and Daniel H. Garrison<sup>1,3</sup>, <sup>1</sup>ARES, code KR, NASA, Johnson Space Center, Houston, TX 77058, <sup>2</sup>NASA Postdoctoral Program fellow, <sup>3</sup>Lockheed Martin, P. O. Box 58561, Houston, TX.**  
[Jisun.park-1@nasa.gov](mailto:Jisun.park-1@nasa.gov)

**Basaltic Shergottite DaG (Dar al Gani) 476.** was found in Libya in 1998, together with paired samples DaG 489/670/735/876/975. It has mega-size (max. 5 mm) olivine phenocrysts (10–24 %) pyroxene (54–65 %), fine-grained maskelynite (12–17%), and mesostasis, along with minor phases of opaques, impact melt glass, carbonates etc. (1–4). DaG-476 gives evidence of terrestrial weathering (1,2,3,5,6,7). The Sm-Nd age of DaG 476 was reported as  $474 \pm 11$  Ma (7, 8). Garrison and Bogard (9) suggested several mixed trapped Ar components with older age of  $\sim 474$  Ma. Walton et al. (10) reported 1427 Ma with big errors.

**Methods.** Our DaG 476 sample was kindly provided by K. Nagao, University of Tokyo. A 100–200 mesh fraction was separated into minerals by C.-Y. Shih and J. Park. Plagioclase glass (8.77 mg) was concentrated using heavy liquid with density cut at  $<2.85\text{g/cm}^3$ . Two different pyroxenes, labeled ‘Lite’ (Mg-rich pyroxene, 27.4 mg) and ‘Dark’ (Fe-rich pyroxene, 29.82 mg), were obtained with density cuts at  $2.96 < \rho < 3.32\text{ g/cm}^3$  and  $3.32 < \rho < 3.45\text{ g/cm}^3$ , respectively. All separated samples were irradiated at the University of Missouri Research Reactor and the Ar isotopic compositions were measured on a VG-3600 mass spectrometer at NASA-JSC. DaG 476 whole rock sample had been reported by Garrison and Bogard (9).

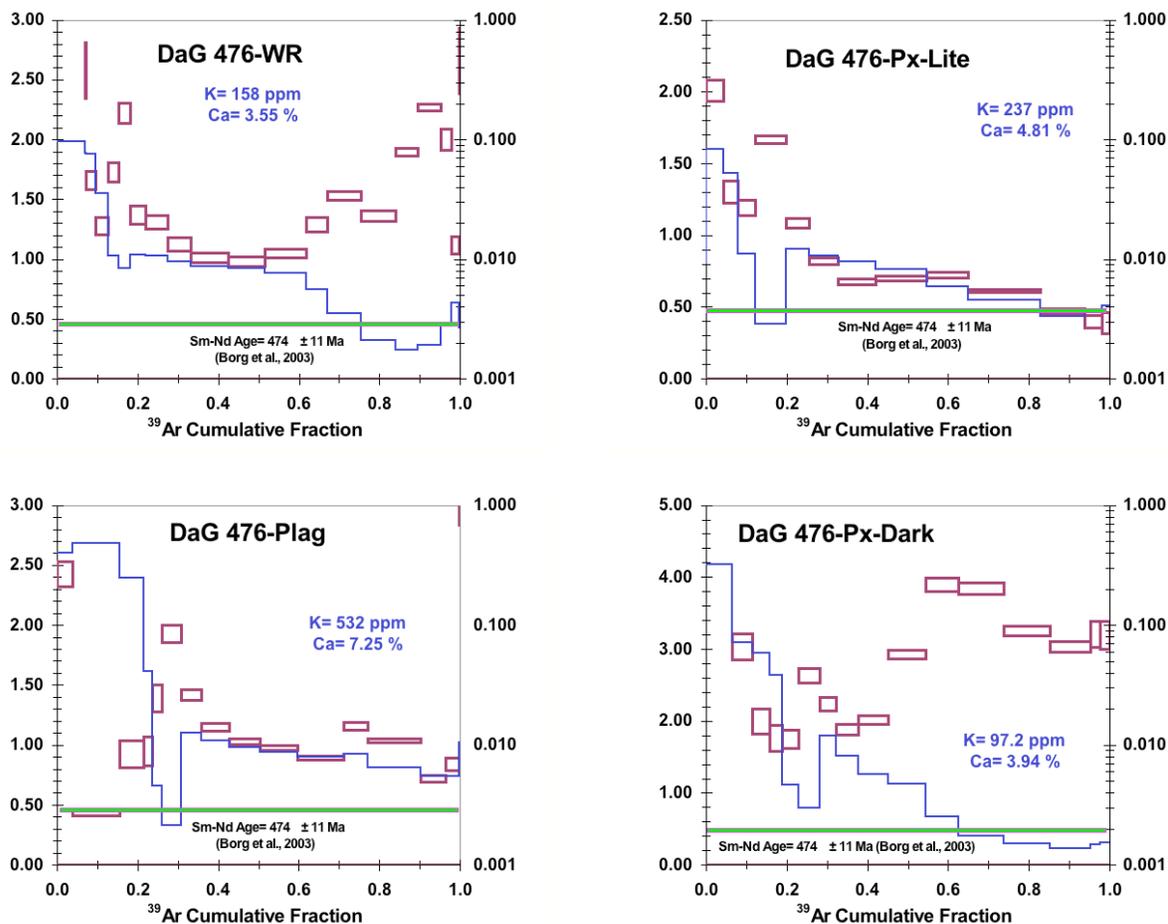
**Apparent Ar-Ar Ages & Excess  $^{40}\text{Ar}$ .**  $^{39}\text{Ar}$ - $^{40}\text{Ar}$  ages for whole rock (WR) and mineral separates of maskelynitized plagioclase (Plag), both pyroxenes, Px-Lite and Px-Dark are shown in Figures 1–4. All apparent Ar-Ar ages are older than the Sm-Nd age of  $474 \pm 11$  Ma (7). Ar-Ar age spectra suggest variable mixtures of terrestrial Ar (at low extraction temperatures), radiogenic Ar, and a trapped Ar component. A  $^{40}\text{Ar}/^{36}\text{Ar}_{\text{cos}}$  vs.  $^{39}\text{Ar}/^{36}\text{Ar}_{\text{cos}}$  isochron plot of DaG 476 WR, with 7–62% of  $^{39}\text{Ar}$  release, gives an age of 882 Ma. A  $^{40}\text{Ar}/^{36}\text{Ar}_{\text{cos}}$  vs.  $^{39}\text{Ar}/^{36}\text{Ar}_{\text{cos}}$  isochron plot of DaG

476 Plag with 36–71 % of  $^{39}\text{Ar}$  release, gives an age of  $619 \pm 31$  Ma. The isochron age of DaG 476-Px-Lite, with 25–83% of  $^{39}\text{Ar}$  release, gives an age of  $408 \pm 172$  Ma. DaG-Px-Dark does not define an isochron.

We calculated the amounts of excess  $^{40}\text{Ar}$  in each DaG-476 sample by subtracting from total  $^{40}\text{Ar}$  that amount of  $^{40}\text{Ar}$  that would have accumulated over 474 Myr. Table 1 shows that these excess  $^{40}\text{Ar}$  concentrations range over  $5\text{--}22.4 \times 10^{-7}\text{ cm}^3/\text{g}$ . This range of excess  $^{40}\text{Ar}$  concentrations is identical to that determined in Zagami and several other shergottites (11, 12). These authors concluded that excess  $^{40}\text{Ar}$  in many shergottites was inherited from the magma, and is not a component implanted from the martian atmosphere, as occurs in shergottite impact glass.

**Terrestrial Weathering.** The terrestrial age of DaG-476, based on  $^{81}\text{Kr}$ -Kr analysis, was reported as 140 Ma (13), which probably accounts for the observed evidence of significant terrestrial weathering (1,2,3,5,6,7). We found that 60% of  $^{40}\text{Ar}$ , 50% of  $^{38}\text{Ar}$  and 73% of  $^{36}\text{Ar}$  in DaG 476-WR were extracted at the first temperature fraction. For DaG 476-Plag and DaG 476-Lite about 10% of  $^{40}\text{Ar}$ ,  $^{38}\text{Ar}$ ,  $^{36}\text{Ar}$  and less than 1% of  $^{40}\text{Ar}$ ,  $^{38}\text{Ar}$ ,  $^{36}\text{Ar}$ , were extracted at the first temperature fraction, respectively. 57% of  $^{40}\text{Ar}$ , 16% of  $^{38}\text{Ar}$  and 74% of  $^{36}\text{Ar}$  in DaG 476-Px-Dark were extracted at the first temperature fraction of  $300^\circ\text{C}$ . Consequently, the DaG 476-Px-Dark which composed of Fe rich pyroxene was easily weathered by terrestrial contamination, but Mg-rich pyroxene was more resistant to weathering.

**Conclusions.** Ar-Ar analyses of DaG 476 whole rock, plagioclase glass, and two pyroxene separates give apparent Ar-Ar ages much older than the reported Sm-Nd age (7). These DaG-476 samples show identical concentrations of excess, trapped Ar to several other shergottites. We conclude that the origin of much of the excess



Figures 1-4.  $^{39}\text{Ar}$ - $^{40}\text{Ar}$  ages (rectangles, left axis) and K/Ca ratios (stepped line, right axis) for stepwise temperature extractions of DaG476 separates. The pink color-line indicates the Sm-Nd age reported by (7).

$^{40}\text{Ar}$  in shergottites was from the shergottite magma (11, 12). But, DaG 476 samples, especially Fe-rich pyroxene, show greater effects of terrestrial weathering compared to most shergottites.

Table 1. Excess  $^{40}\text{Ar}$  in Zagami Mineral Separates (units cc/g and percent of total  $^{40}\text{Ar}$ ).

	$^{40}\text{Ar}$ excess (cc/g)	$^{40}\text{Ar}$ excess (%)
DaG476-WR	9.64E-07	75.4
DaG476-Plag	1.40E-06	59.7
DaG476-Px-Lite	5.05E-07	53.4
DaG476-Px-Dark	2.24E-06	92.5

#### References.

- (1) Zipfel et al., 2000, *MAPS* 35, 95,
- (2) Mikouchi et al., 2001, *MAPS* 36, 531,
- (3) Folco et al., 2000, *MAPS* 35, 827,
- (4) Wadhwa et al., 2001, *MAPS* 36, 195,
- (5) Crozaz & Wadhwa, 2001, *GCA* 65, 971,
- (6) Greshake & Stoeffler, 1999, *LPSC* 30, #1377,
- (7) Borg et al., 2003, *GCA* 67, 3519,
- (8) Borg et al., 2001. *LPSC* 32, #1144,
- (9) Garrison & Bogard, 2001, *Annal Meteorit. Soc. Meeting* 64, #5087,
- (10) Walton et al., 2006, *GCA* 71, 497,
- (11) Bogard & Park, 2008a, *MAPS*, in press.
- (12) Bogard & Park, 2008b, *LPSC* 39, this volume,
- (13) Park et al., 2003, *LPSC* 34, #1213,