

**SHEAR HEATING AND THE ORBITAL EVOLUTION OF ENCELADUS.** F. Nimmo, *Dept. Earth & Planetary Sciences, U.C. Santa Cruz, Santa Cruz CA 95064 (fnimmo@es.ucsc.edu).*

Volumetric dissipation within satellites is typically assumed to be viscoelastic and is calculated using the Love number  $k_2$  and quality factor  $Q$  [1]. However, observations of the plumes and “tiger stripes” at Enceladus [2] have led to suggestions that tidal dissipation may primarily be accomplished by shear-heating along individual faults [3]. This style of dissipation has a different functional form to the usual viscoelastic formula. In particular, the eccentricity damping timescale will differ, which may have important implications if Enceladus’s eccentricity is time-dependent [4,5].

We assume that tidal stresses  $\sigma$  are capable of moving faults to a depth  $d$ . Here  $d = \sigma / \rho g f$  where the denominator is the overburden pressure resisting fault motion,  $\rho$  is surface density,  $g$  is gravity and  $f$  is the coefficient of friction. The tidal strain rate  $\dot{\epsilon}$  depends on orbital parameters and the Love number  $h_2$  and is linearly proportional to  $\sigma$ . For active faults of total length  $L_{tot}$  and separated by an average distance  $w$ , the total frictional heating on the faults  $\dot{W}$  depends on the fault depth and mean shear velocity, and is given by

$$\dot{W} = \frac{\dot{\epsilon}^3 L_{tot} w \mu'^2}{2 \rho g f n^2} \quad (1)$$

where  $\mu'$  (units: Pa) relates stress to strain and  $n$  is the mean motion of the satellite. Note that here we are neglecting the contribution from viscous shear heating, which is generally comparable to the frictional contribution [6], and any volumetric dissipation in the interior.

Substituting in for the tidal strain rate [7] we obtain

$$\dot{W} = n^7 e^3 h_2^3 \frac{L_{tot} w \mu'^2 c^3}{2 \rho g f G^3 \bar{\rho}^3} \quad (2)$$

where  $c$  is a constant (1.55),  $G$  is the universal gravitational constant and  $\bar{\rho}$  is the bulk density. This expression has a different functional form to the standard expression for tidal dissipation, which depends on  $k_2/Q$ ,  $e^2$  and  $n^5$  [1].

In the absence of external torques, the rate of change of semi-major axis  $a$  and eccentricity  $e$  of a satellite are linked. Dissipation in the satellite results in a reduction in  $a$ ; conservation of angular momentum then allows the change in eccentricity  $e$  to be determined:

$$\frac{de}{dt} = -\frac{1-e^2}{e} \frac{a}{G m_p m_s} \dot{W} \quad (3)$$

where  $m_p$  and  $m_s$  are the mass of the primary and satellite, respectively. This expression may be integrated if the functional form of  $\dot{W}$  is known.

Figure 1 shows the effect of applying both the conventional dissipation (a) and the shear-heating (b) calculations to a body with present-day characteristics appropriate to Enceladus. The main difference of the latter approach is that the

effective eccentricity damping timescale is larger by 70%. The dissipation thus falls off more slowly as a function of time (but more rapidly as a function of eccentricity). A lower rate of eccentricity reduction makes it more likely that oscillatory behaviour [4] can occur.

#### References

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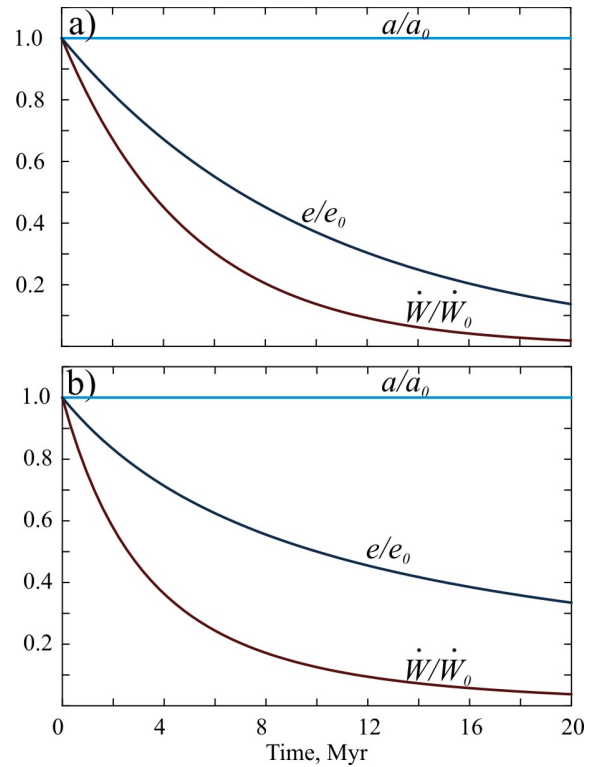


Figure 1: Evolution of  $a, e$  and heat production rate  $\dot{W}$  with time, calculated using equation (3). Values are normalized to initial values ( $0.238 \times 10^6$  km, 0.0045 and 1 GW, respectively) appropriate to present-day Enceladus. Panel a) gives evolution using conventional tidal dissipation formula; panel b) gives evolution using shear-heating formula (equation 2). Effective eccentricity damping timescales are 10.2 Myr and 17.5 Myr, respectively. Here it is assumed that the characteristics of the satellite (especially  $h_2$ ) do not vary in time.