

**NITROGEN ON EARLY TITAN.** A.A. Berezhnoy, Sternberg Astronomical Institute, Moscow State University, Universitetskij pr., 13, Moscow, Russia, Email: [ber@sai.msu.ru](mailto:ber@sai.msu.ru)

**Introduction:** The origin of volatiles on Titan can be understood based on isotopic abundances of N, C, and H in Titan's atmosphere. The initial mass of volatiles on early Titan is a very important parameter for constructing the early history of the system of Saturnian satellites. In this work we try to understand the early history of Titan's main N-containing compounds, N<sub>2</sub> and NH<sub>3</sub>.

**Model of calculation of isotopic composition of nitrogen on Titan:** The nitrogen in Titan's atmosphere is assumed to be escaping from the exobase level. Below the exobase and above the homopause diffusive separation of species with different molar masses occurs and leads to enrichment of lighter species at the exobase level. Dissociative mass fractionation occurs due to differences between the rate constants of removal processes such as electron impact, solar wind bombardment, and photodissociation and between the velocity distributions of photolysis products for species containing different isotopes. The initial mass of Titan's atmosphere  $M_0$  is calculated from the current atmospheric mass  $M_{pr}$  and  $^{14}\text{N}/^{15}\text{N}$  ratio assuming that this ratio was changed from its original value mainly by two processes: 1) dissociative fractionation  $y_d$  and escape, which depends on the probabilities of removal processes for different isotopes ( $y_{d1}$  factor), escape velocity, and velocity distribution of N-containing species ( $y_{d2}$  factor) produced by photolysis, electron impact, and solar wind bombardment; and 2) diffusive fractionation  $y_f$ , which depends on the heights of the homopause and exobase and the temperature of the upper atmosphere.

Based on the available  $^{14}\text{N}/^{15}\text{N}$  ratio in HCN on Titan, about 60 (Hidayt and Marten [1]), the initial mass of Titan's atmosphere is estimated in wide range between 10-1000  $M_{pr}$  [2]. In this work the exobase radius  $R_{ex}$  as measured from the center of Titan is taken as 4175 km [3]. The homopause level  $R_{hp}$  has been constrained from Voyager UV methane spectra to lie at 3560 to 3700 km [3]. Here it is assumed that homopause and exobase levels on early Titan were the same as they are today.

Measurements made by the Huygens probe show that  $^{14}\text{N}/^{15}\text{N}$  value in N<sub>2</sub> on Titan is  $183 \pm 5$  [4]. This N<sub>2</sub> ratio is representative of nitrogen in the entire atmosphere because N<sub>2</sub> is the main N-containing compound. It is also 1.5 times lower than terrestrial value, which is assumed to be the same as the primordial value for early Titan. Using enrichment factor  $F = 1.5$  instead of 4 [2] Niemann et al. [4] estimated the initial atmospheric mass as 2 - 10  $M_{pr}$ . Here we re-examine this estimate by study of N<sub>2</sub> and NH<sub>3</sub> photolysis on Titan.

**Fractionation of nitrogen isotopes during photolysis of N-containing molecules:** In [2] it was

assumed that N<sub>2</sub> was the main N-containing compound in the early Titan's atmosphere just as it is today. However, calculated values of N loss rate from Titan [5] are low and Titan's N<sub>2</sub> is therefore thought to have been stable over billions years. Loss of several current atmospheric masses required for explanation of current  $^{14}\text{N}/^{15}\text{N}$  ratio can occur during short strong flashes of intensity of the solar wind at the early period of history of the Solar system [6]. Dissociative fractionation on early Titan is difficult to estimate because the energy distributions of solar photons, solar wind particles, and electrons are unknown at that time.

The calculated mole fractions of CN and HCN at the exobase level of Titan are about  $10^{-3}$  [7]. The photolysis rates calculated here for CN, HCN, and N<sub>2</sub> show that most of the N loss from Titan today comes from N<sub>2</sub>. However, based on P-T conditions of stability of N-containing clathrates, on early Titan NH<sub>3</sub> is considered as the main N-containing compound. An early NH<sub>3</sub>-rich atmosphere can be quickly converted into N<sub>2</sub>-rich atmosphere because present-day NH<sub>3</sub> photolysis rate is very high [8].

The dissociative fractionation  $y_d$  of N<sub>2</sub> was allowed to vary from 0 to 0.6 in paper [2]. Let us estimate this value more precisely. The present-day fractionation of N isotopes depends mainly on the ratio of the relative removal rates of  $^{14}\text{N}^{15}\text{N}$  and  $^{14}\text{N}_2$ . The main reaction controlling N loss from current Titan is electron-impact dissociation of N<sub>2</sub> but, unfortunately, the cross section of this reaction is known only for  $^{14}\text{N}_2$  [9], and not for  $^{14}\text{N}^{15}\text{N}$ . The photolysis rate of  $^{15}\text{NH}_3$  is 12 % higher than that of  $^{14}\text{NH}_3$  [10]. One can expect a higher photolysis rate and a lower lifetime against electron-impact dissociation for  $^{14}\text{N}^{15}\text{N}$  than for  $^{14}\text{N}_2$ . Electron-impact dissociation can even lead to negative values of the factor  $y_{d1}$ . Based on the discussion above, the factor  $y_{d1}$  is estimated in the range of -0.2 – 0.2. Higher values are unlikely because there are no known processes that produce nitrogen with typical velocities that approach the escape velocity. Lower values are also unlikely because we do not expect a significant difference (more than 20 %) in rates of photolysis, electron-impact dissociation and dissociative recombination of  $^{14}\text{N}$  and  $^{15}\text{N}$ -containing species.

Additional dissociation fractionation characterised by the factor  $y_{d2}$  occurs when the typical velocity of N-containing photolysis products is close to the escape velocity from Titan. The velocity distribution of N and NH<sub>2</sub> produced via N<sub>2</sub> and NH<sub>3</sub> photolysis is estimated from the present-day solar photon flux [8] and the wavelength-dependent photolysis rates of N-containing species [11]. We assume that the kinetic energy of molecules before

photolysis is negligible in comparison to the kinetic energy of the products of this reaction. The kinetic energy of the photolysis products is estimated from the laws of conservation of energy and angular momentum with the added but experimentally based assumption [12] that 25 % of all excess energy (the difference between the energy of photon and the threshold energy for photodissociation) is converted to kinetic energy. The velocity distribution of photolysis products is non-Maxwellian; most probable velocities of  $\text{NH}_2$  produced via  $\text{NH}_3$  photolysis are 600 - 800 m/s. The factor  $y_{d2}$  is significantly lower than the values of the present-day diffusion fractionation  $y_f$ , about 0.25 - 0.4 [2]. Specifically,  $y_{d2}$  is close to zero for  $\text{N}_2$  and  $\text{NH}_3$  photolysis, also for different reasons. The typical velocity of N atoms formed by  $\text{N}_2$  photolysis is about 4 km/s, a value is significantly higher than the escape velocity from the current exobase level, 2.07 km/s, while photolysis-generated  $\text{NH}_2$  radicals remain on Titan. The  $\text{NH}_3$  photolysis does not change significantly the isotopic composition of Titan's atmosphere because nitrogen is remaining in the atmosphere after  $\text{NH}_3$  photolysis and  $\text{NH}_3$  is minor reservoir of hydrogen on Titan.

In paper [2]'s treatment of early Titan it was assumed that the temperature, homopause and exobase levels of Titan's atmosphere remained constant though time. Changing of these parameters gives additional uncertainty to estimates of the initial atmospheric mass. For example, at  $R_{\text{ex}} = 4175$  km,  $R_{\text{hp}} = 3700$  km,  $y_d = 0$ , and  $F = 1.5$  the initial mass  $M_0$  is equal to 7, 14, 26  $M_{\text{pr}}$  at temperatures of upper atmosphere equal to 150, 200, and 250 K, respectively. However, if we assume that the exobase and homopause altitudes increase linearly with increasing temperature of the upper atmosphere, then  $M_0$  is independent of the temperature of the upper atmosphere.

Further, the presence of  $\text{NH}_3$  may have influenced of  $R_{\text{ex}}$  and  $R_{\text{hp}}$ . Specifically, the scale height, which depends inversely on average molar mass, would have been 1.7 larger in an  $\text{NH}_3$ -rich atmosphere than in a  $\text{N}_2$ -rich one. A 1.5-2  $\times$  increase of the exobase altitude leads to a decrease of the escape velocity from Titan to 1750 - 1900 m/s, but it does not increase significantly the rate of N loss from Titan and  $y_{d2}$  value. The fractionation of nitrogen isotopes is controlled by diffusive separation of isotopes between homopause and exobase on the present-day Titan. Discussed above estimates of the  $y_{d1}$ ,  $y_{d2}$ , and  $y_f$  factors corresponds to the value of the initial mass of the Titan's atmosphere of about 3-20  $M_{\text{pr}}$ .

**Fractionation of hydrogen isotopes during photolysis of  $\text{NH}_3$ :** The observed D/H value on Titan cannot be explained only by isotopic fractionation during  $\text{CH}_4$  photolysis. The D/H ratio on early Titan is estimated to be 3.6 - 4.3 $\times$  and 2.7 - 3.6 $\times$  the

protosolar values for the cases of presence of  $\text{CH}_4$  reservoir on Titan during 0.6 and 4.5 Gyr, respectively [13]. The loss of  $\text{NH}_3$  from early Titan may have increased the D/H ratio by an additional factor of 1.1. For these calculations we take the ratio of  $\text{NH}_2\text{D}$  and  $\text{NH}_3$  photolysis rates equal to be 0.71 [14] and the N/C ratio in the protosaturnian nebula equal to the solar value, 0.275 [15]. Such small enrichment factor, about 1.1, disagrees with assumption of low D/H value in protosaturnian nebula equal to that in protosolar nebula. Thus, primordial D/H value on Titan is higher than corresponding protosolar value [13]. Primordial  $^{14}\text{N}/^{15}\text{N}$  ratio on Titan is also suggested to be lower than corresponding protosolar value [16].

Comet-like objects as sources of Titan's volatiles can lead to additional  $^{15}\text{N}$  enrichment of Titan's atmosphere because the average  $^{14}\text{N}/^{15}\text{N}$  ratio in cometary CN, about 140, is 1.9 times less than the terrestrial value [17]. If the primordial  $^{14}\text{N}/^{15}\text{N}$  ratio on Titan was lower than the corresponding protosolar ratio then our estimation of the initial mass of Titan's atmosphere can be considered as an upper limit.

**Conclusions:** The mechanisms of losing nitrogen from early Titan's atmosphere are still unknown. An early  $\text{NH}_3$ -dominated atmosphere can be quickly converted into the present-day  $\text{N}_2$ -rich atmosphere without significant changes in the isotopic composition of N and H. The value of the dissociative fractionation factor  $y_d$  is estimated for present-day photolysis, it can vary in wide range of -0.2 to +0.2. This range leads to significant uncertainty in estimates of the upper limit of the initial atmospheric mass of Titan, about 3-20  $M_{\text{pr}}$ . Additional fractionation of N isotopes in the protosaturnian nebula may be important.

**Acknowledgments:** This work was supported by RFBR grant 08-05-01070-a. The author thanks Ashfold M.N.R., Herzog G., and Dorofeeva V.A. for helpful comments and suggestions.

**References:** [1] Hidayat, T., Marten, A. (1998) *Annales Geophysicae* 16 (Suppl. III), C998. [2] Lunine J.L. et al. (1999) *Plan. Sp. Sci.*, 47, 1291-1303. [3] Strobel D.F. et al. (1992) *Icarus*, 100, 512-526. [4] Niemann H.B. et al. (2005) *Nature*, 438, 779-784. [5] Shematovich V.I. et al. (2003) *JGR*, 108, doi:10.1029/2003JE002094 [6] Lammer H., Bauer S.J. (2003) *Space Science Reviews*, 106, 281-291. [7] Yung Y.L. et al. (1984) *Astrophys. J. Supplem. Series*, 55, 465-506. [8] Huebner W. F. et al. (1992) *Astrophys. Space Sci.*, 195, 1-289, 291-294. [9] Cosby P.C. (1993) *J. Chem. Phys.*, 98, 9544-9553. [10] Liang M.-C. et al. (2007) *ApJ*, 657, L117-120. [11] <http://amop.space.swri.edu> [12] Ashfold M.N.R. et al. (1997) *Phil. Trans. R. Soc. London*, 355, 1659-1676. [13] Cordier D. et al. (2008) *ApJ*, 689, L61-L64. [14] Cheng B.-M. et al. (2006) *ApJ*, 647, 1535-1542. [15] Lodders K. (2003) *ApJ*, 591, 1220-1247. [16] Mousis O. et al. (2002) *Icarus*, 159, 156-165. [17] Schulz R. et al. (2008) *Planetary and Space Science*, 56, 1713-1718.