

**NOBLE GAS RECORD OF THE ANOMALOUS CM DHOFAR 1434.** R. Bartoschewitz<sup>1</sup>, U. Ott<sup>2</sup>, and S. Herrmann<sup>2</sup>. <sup>1</sup>Bartoschewitz Meteorite Laboratory, Lehmweg 53, D-38518 Gifhorn, Germany. E-mail: bartoschewitz.meteorite-lab@t-online.de. <sup>2</sup>Abt. Biogeochemie, Max-Planck-Institut für Chemie, Joh.-J.-Becher-Weg 27, D-55128 Mainz, Germany

**Introduction:** We report noble gases for Dhofar 1434 that has been classified as anomalous CM chondrite [1]. It was discovered about 600 m west of Dhofar 225 that is classified as anomalous CM, too [2]. Petrological and mineralogical data of Dhofar 1434 are very close to Dhofar 225 [3]; and both meteorites seem to be paired. The O-isotopes plot on the far extension of CM2 chondrites and have affinities to the anomalous CM chondrites Belgica-7904, Y-82162, and Y-86720 [3].

**Samples and Experimental:** Noble gases were analyzed by standard analytical techniques (e.g., [4]) in two fragments (sample I: 8.54 mg; sample II: 14.7 mg). Gas extraction was at 600°C, 1000°C, and 1800°C.

**Results:** Data for He and Ne are summarized in Table 1, those for Ar, Kr and Xe in Table 2.

#	<sup>3</sup> He	<sup>4</sup> He	<sup>22</sup> Ne	<sup>20</sup> Ne/ <sup>22</sup> Ne	<sup>21</sup> Ne/ <sup>22</sup> Ne
I	0.5663 ±0.0134	1915 ±60	1.110 ±0.027	5.771 ±0.043	0.2346 ±0.0031
II	0.5100 ±0.0087	1760 ±44	1.074 ±0.021	5.562 ±0.034	0.2536 ±0.0025

Table 1. He and Ne results (abundances in 10<sup>-8</sup> cc/g).

#	<sup>36</sup> Ar	<sup>38</sup> Ar/ <sup>36</sup> Ar	<sup>40</sup> Ar	<sup>84</sup> Kr	<sup>132</sup> Xe
I	7.799 ±0.172	0.1871 ±0.0008	75.0 ±5.2	0.0891 ±0.0032	0.0942 ±0.0023
II	7.529 ±0.136	0.1862 ±0.0008	80.1 ±3.2	0.0792 ±0.0026	0.0918 ±0.0017

Table 2. Ar, Kr and Xe results (abundances in 10<sup>-8</sup> cc/g).

**Helium and neon.** Neon data are shown in the 3-isotope plot of Fig. 1. The dominant trapped component is Ne-HL (Ne-A2) carried by presolar diamond [5, 6], with indications for a small amount of Ne with higher <sup>20</sup>Ne/<sup>22</sup>Ne (air, P3, solar) released at the lowest temperature. Evident is also the presence of presolar SiC as carrier for Ne-E (Ne-G) indicated by the low <sup>20</sup>Ne/<sup>22</sup>Ne [6, 7] at 1800 °C. With a trapped <sup>20</sup>Ne abundance of ~6x10<sup>-8</sup> cc/g a significant amount – if not virtually all – of measured <sup>4</sup>He must be of trapped origin. In fact, for a trapped <sup>4</sup>He/<sup>20</sup>Ne ratio of ~600 (as in primitive HL gases; [6, 8]) the expected amount of trapped <sup>4</sup>He is about twice that observed; For the <sup>4</sup>He/<sup>20</sup>Ne ratio of “processed HL” gases [8] of ~ 290, on the other hand, the expected trapped <sup>4</sup>He agrees

well with the observed amount, which then leaves little room for radionogenic He. Almost half of <sup>3</sup>He would then be trapped, with corresponding uncertainties about the amount of cosmogenic He (see below).

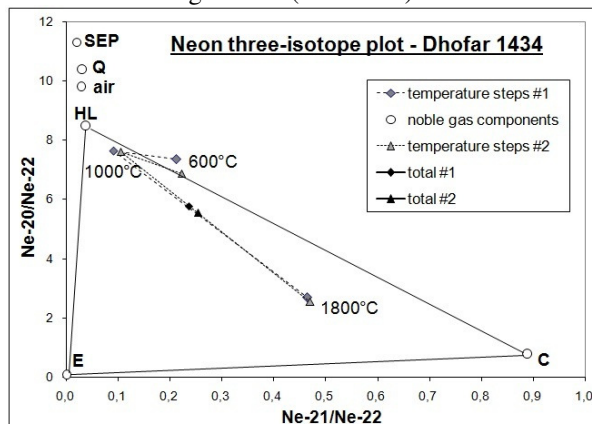


Fig. 1. Neon 3-isotope plot. The solid line spans the triangle defined by Ne-E, Ne-HL and a typical composition for cosmogenic Ne (C). The dashed line traces the stepped analysis. Reference compositions are from [6, 9]. SEP refers to the now discredited SEP component [10].

**Argon, krypton, xenon.** Ar, Kr and Xe are dominated by Q-gases [6, 11]. As shown in Fig. 2, a small contribution from adsorbed air is evident in the first release step, while HL-Xe carried by presolar nanodiamond [5, 6] makes a significant contribution (~25% of <sup>132</sup>Xe) at intermediate temperature (1000 °C).

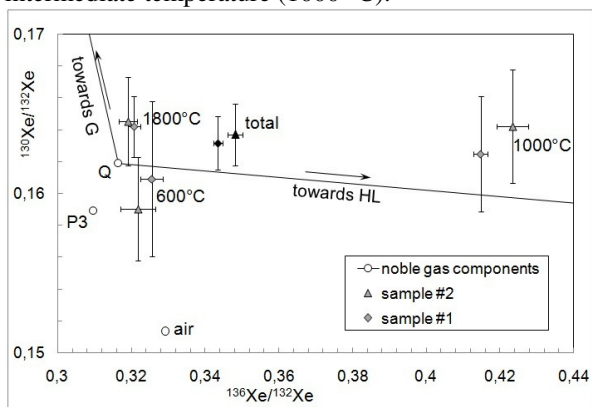


Fig. 2. Xenon 3-isotope plot of <sup>130</sup>Xe/<sup>132</sup>Xe vs. <sup>136</sup>Xe/<sup>132</sup>Xe. The lines are mixing lines between Xe-Q and a) HL-Xe carried by presolar diamond off scale to the right and b) Xe-G carried by presolar SiC towards the upper left. Reference compositions from [6].

*Abundance of presolar phases.* The isotopic excursions in Ne (Fig. 1) and Xe (Fig. 2) allow an estimate of the abundances of presolar diamond and SiC. Using the approach of [12, 13] and a “constant” Ne-E (= Ne-G) concentration in SiC of  $16500 \times 10^8$  cc/g [13], from the 1800 °C steps an abundance of ~6 ppm is inferred. A similar estimate for presolar diamond, based on  $^{132}\text{Xe}$ -HL in the 1000°C steps of  $\sim 6 \times 10^{-11}$  cc/g and a typical  $^{132}\text{Xe}$ -HL abundance in CM diamonds of  $1.5 \times 10^{-7}$  cc/g [8, 13] results in an estimated diamond abundance of ~400 ppm.

*Ages.* Assuming that Dho 1434 is paired with Dho 225 we used the chemical composition from [3] for age calculations. Nominal K-Ar ages based on the  $^{40}\text{Ar}$  abundances in Table 2 (probably containing a significant air contribution) are 0.23 and 0.24 Ga. With a probably dominant contribution to  $^4\text{He}$  from trapped He (see above), it is not possible to calculate a U/Th-He age. Cosmic ray exposure (CRE) ages that we have calculated following [14] suffer from the problem that the abundance of trapped Ne is too high for a reliable determination of the shielding parameter  $(^{22}\text{Ne}/^{21}\text{Ne})_{\text{cos}}$ . Assuming “average shielding” (shielding parameter = 1.11) leads to CRE ages from  $^{21}\text{Ne}$  of 0.74 and 0.79 Ma. Corresponding ages based on  $^3\text{He}$  and the unrealistic assumption that all of  $^3\text{He}$  is cosmogenic (see above) are 0.35 and 0.32 Ma, reflecting loss of at least half of cosmogenic helium.

**References:** [1] Weisberg M. K. et al. (2009) *Meteoritics & Planet. Sci.*, 44, 429-462. [2] Russell S. S. et al. (2002) *Meteoritics & Planet. Sci.*, 37, A157-A184. [3] Ivanova M. A. et al. (2010) *Meteoritics & Planet. Sci.*, 45, 1108-1123. [4] Schwenzer S. P. et al. (2007) *Meteoritics & Planet. Sci.*, 42, 387-412. [5] Lewis R. S. (1987) *Nature*, 326, 160-162. [6] Ott U. (2002) in *Reviews in Mineralogy & Geochemistry Vol. 47*, 71-100. [7] Tang M. et al. (1988) *Geochim. Cosmochim. Acta*, 52, 1235-1244. [8] Huss G. R. and Lewis R. S. (1994) *Meteoritics* 29, 811-829. [9] Wieler R. (2002) in *Reviews in Mineralogy & Geochemistry Vol. 47*, 125-170. [10] Wieler R. (2007) *Chemical Geology* 244, 382-390. [11] Lewis R. S. et al. (1975) *Science*, 190, 1251-1262. [12] Huss G. R. et al. (1995) *Geochim. Cosmochim. Acta*, 59, 115-160. [13] Huss G. R. et al. (2003) *Geochim. Cosmochim. Acta*, 67, 4823-4848. [14] Eugster O. (1988) *Geochim. Cosmochim. Acta*, 52, 1649-1662.