

OBJECTIVES: Understanding the environments and isotope reservoirs of R chondrite forming regions.

Characteristics of R chondrites [1-7]

- FeO-rich olivine (Fa₃₈) in equilibrated lithologies
- High bulk $\Delta^{17}\text{O}$ ($=\delta^{17}\text{O}-0.52\times\delta^{18}\text{O}$) values ($\sim 2.7\text{‰}$)
- A wide range of $\Delta^{17}\text{O}$ in chondrules (-4‰ to $+3\text{‰}$)

Oxidized nebula setting with ¹⁶O-depleted oxygen isotope reservoirs?

Search for least equilibrated R3 clasts

Most R chondrites are of lower petrologic types (R3 and R4), though unequilibrated R chondrites with subtypes lower than 3.6 are not known. Many R chondrites are regolith breccias containing R3-6 clasts. Unequilibrated R3 clasts (<3.5) have been observed in brecciated R chondrites [4,5]. In order to identify primary oxygen isotope signatures in R chondrite chondrules, we need to search for R3 clasts that experienced minimal parent body metamorphism, equivalent to lowest subtypes (≤ 3.2).

We performed detailed petrographic descriptions of chondrules in a highly unequilibrated R3 clast from Northwest Africa (NWA) 753 (R3-5) chondrite (Figs. 1,2). The goal of the study is to determine if the clast is suitable for detailed SIMS oxygen isotope analyses.

- NWA 753 "Fragment IX" (8 mm × 4 mm)
- We found 69 chondrules >200 μm in size
- BSE imaging (Hitachi S-3400N SEM)
- SEM-EDX analysis for identification of phases
- Electron Microprobe analyses (Cameca SX51)



Fig. 1. Transmitted light photograph of the clast NWA 753-IX in thin section. The clast (marked by red dashed line) stands out by the dark fine-grained matrix. The scale bar is 1 mm.

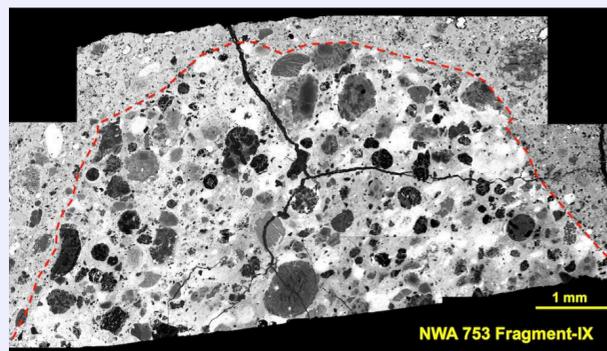


Fig. 2. Backscattered electron image of the clast NWA 753-IX. The clast stands out by the high contrast between chondrules and fine-grained matrix. The scale bar is 1 mm.

Chondrules in R3 clast NWA 753-IX

Most chondrules show porphyritic textures, while minor non-porphyritic chondrules and two Al-rich chondrules [8] are also identified (Table 1, Fig. 3). The proportion of different types of chondrules are very similar to LL3 chondrites [9], especially with respect to similar abundances of type I and II chondrules, while type II chondrules are less abundant in carbonaceous chondrites ($\sim 25\%$ in CO3 and $\leq 5\%$ in CR3 [10-11]). All non-porphyritic chondrules are FeO-rich.

Mesostasis of chondrules are mostly altered compared to unaltered mesostasis seen in type 3.0 chondrites (such as Semarkona [9]) and often include areas with high Na and Cl contents (Fig. 4).

Table 1. Classification of chondrules in NWA 753-IX

Porphyritic (N=59)		Non-porphyritic (N=10)
Al-rich ($\text{Al}_2\text{O}_3 > 10\%$) = 2		
type I (Mg# ≥ 90) ^{*1}	IA (OI >80%) = 9 IAB (OI 20-80%) = 11 IB (OI <20%) = 9	RP (Radial Pyroxene) = 6 BO (Barred Olivine) = 3
type II (Mg# <90) ^{*1}	IIA (OI >80%) = 20 IIAB (OI 20-80%) = 7 IIB (OI <20%) = 1	CC (Crypto-crystalline) = 1

^{*1}Mg#: Molar % of [MgO]/([MgO]+FeO)

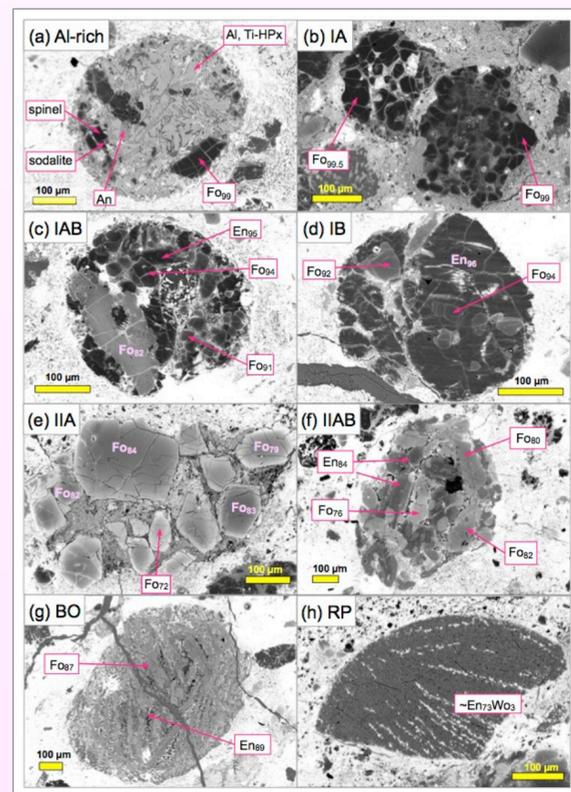


Fig. 3. Major types of chondrules in the R3 clast IX. (a) An Al-rich chondrule, which contains Al-Ti-rich high Ca-pyroxene, and anorthite. (b-f) Ferromagnesian porphyritic chondrules. (g) An FeO-rich barred olivine chondrule. (h) A fragment of radial pyroxene chondrule.

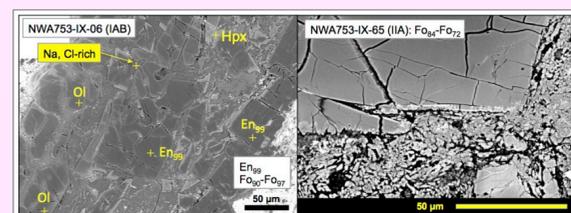


Fig. 4. Alteration textures in the mesostasis of porphyritic chondrules in clast NWA 753-IX. (a) type IAB, (b) type IIA.

Distribution of Chondrule Mg#

The FeO contents of olivine and pyroxene phenocrysts are related to the oxygen fugacity of the chondrule-forming environments [12]. The distribution of Mg# of chondrules in the R3 clast (Fig. 5) is similar to that of LL3, in contrast to carbonaceous chondrites with abundant highly reduced chondrules (Mg#>95). These data indicate that many chondrules in both ordinary chondrite and R chondrite regions formed under more oxidized conditions.

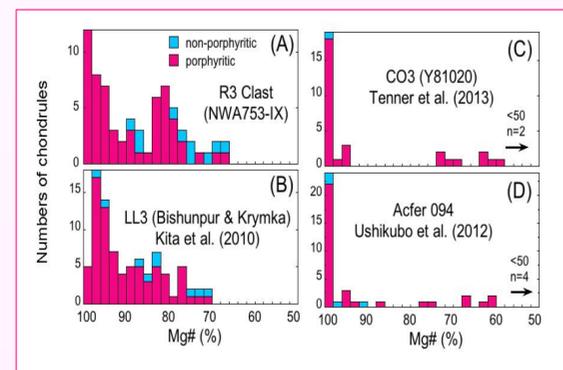


Fig. 5. Histograms of chondrule Mg# in (A) R3 clast NWA 753 IX (this work), (B) LL3 chondrites [9], (C) CO3 [13], and (D) Acfer 094 [14]. The average Mg# of olivine and low-Ca pyroxene phenocrysts were calculated for each chondrule. The Mg# of RP and CC chondrules were also estimated from the average of multiple electron microprobe analyses.

Mineral Chemistry

Most Mg-rich olivine (Mg#>98) in porphyritic chondrules show enrichments of Al_2O_3 and CaO and depletion of MnO when compared to other olivine grains (Fig. 6a-c). This indicates that some type I chondrules in R3 are significantly enriched in refractory elements. They could be related to refractory forsterite grains or forsteritic olivine with ¹⁶O-rich isotope signatures that were reported previously from R3 chondrites [7,15]. The average Cr_2O_3 content of ferroan olivine (Fo<98) is 0.17% with a standard deviation (SD) of 0.11% (Fig. 6d). The average value is much lower than that of Semarkona (LL3.0; $0.50\pm 0.10\%$ [16]).

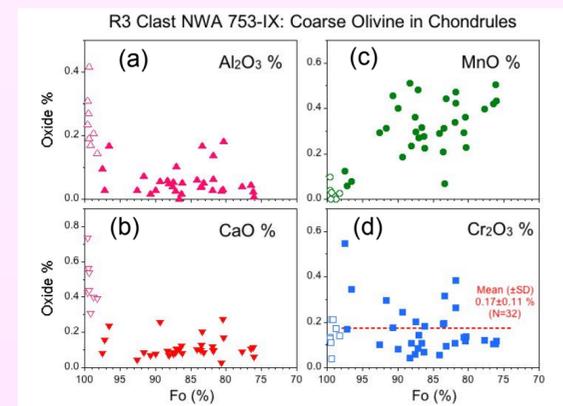


Fig. 6. Minor elements in coarse-grained ($\geq 100 \mu\text{m}$) olivine from porphyritic chondrules in NWA 753 IX. Mg-rich olivines (Fo>98) are shown as open symbols. The detection limits of Al_2O_3 , CaO, MnO, and Cr_2O_3 are 0.04%, 0.03%, 0.09%, and 0.07%, respectively.

References:

- [1] Weisberg M. K. (1991) *GCA*, 55, 2657-2669. [2] Bischoff A. et al. (1994) *Meteoritics*, 29, 264-274. [3] Schulze H. et al. (1994) *Meteoritics*, 29, 275-286. [4] Bischoff A. et al. (2011) *Chemie der Erde*, 71, 101-133. [5] Bischoff A. (2000) *Meteoritics & Planet. Sci.*, 35, 699-706. [6] Greenwood J. P. (2000) *GCA*, 64, 3897-3911. [7] Isa J. (2011) *LPS XLII*, Abstract #2623. [8] Bischoff A. and Keil K. (1984) *GCA*, 48, 693-709. [9] Kita N. T. et al. (2010) *GCA*, 74, 6610-6635. [10] Kunihiro T. et al. (2005) *GCA*, 69, 3831-3840. [11] Tenner T. J. et al. (2012) *LPS XLIII*, Abstract #2127. [12] Ebel D. S. and Grossman L. (2000) *GCA*, 64, 339-366. [13] Tenner T. J. et al. (2013) *GCA*, 102, 226-245. [14] Ushikubo T. (2012) *GCA*, 90, 242-264. [15] Pack A. et al. (2004) *GCA*, 68, 1135-1157. [16] Grossman J. N. and Brearley A. J. (2005) *Meteoritics & Planet. Sci.*, 40, 87-122. [17] Connolly H. C. and Huss G. R. (2010) *GCA*, 74, 2473-2483. [18] Schrader D. L. et al. (2012) *GCA*, 101, 302-327.

Petrologic Subtypes

Grossman and Brearley [16] suggested that the average and SD of Cr_2O_3 wt% in ferroan olivine grains can be used as an index of low petrologic subtypes from 3.00 to 3.15. The data estimated for the R3 clast NWA 753-IX plot close to the subtypes 3.15-3.20 for ordinary chondrites (Fig. 7). By assuming the classification scheme for ordinary chondrites can be also applied to R chondrites the data suggest that clast NWA 753-IX is the least metamorphosed R3 chondritic material compared to other R3 chondrites (≥ 3.6).

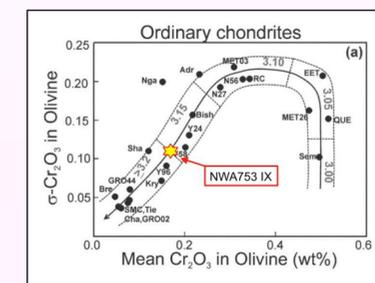


Fig. 7. The average and SD of Cr_2O_3 wt% in olivine (Fo<98) in chondrules from clast NWA 753-IX (Fig. 6d), superimposed on the diagram published by [16].

Future Work: Oxygen isotope Analysis

Group-specific Mg#- $\Delta^{17}\text{O}$ trends: High precision SIMS oxygen isotope studies of chondrules in least equilibrated chondrites (Fig. 8; LL, CO, CR, and Acfer 094 with subtypes ≤ 3.2) revealed group-specific trends between the Mg# and $\Delta^{17}\text{O}$ values of chondrules [e.g., 9, 11, 13-14]. These results provide important clues to identify the variety of isotope reservoirs and their evolution in each chondrule forming region.

Both high and low $\Delta^{17}\text{O}$ values ($+2.7\text{‰}$ and -4.0‰) from Mg-rich chondrules (Mg# ≥ 99) have been reported for chondrules in R3.6 [7], which are important in understanding the origin of ¹⁶O-poor isotope reservoirs in the solar nebula.

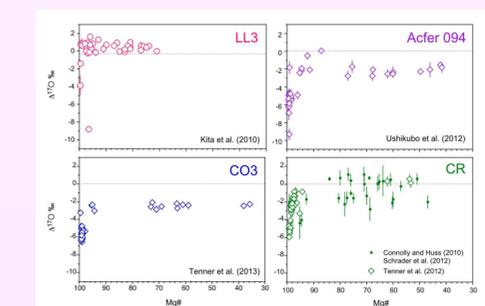


Fig. 8. Relationship between Mg# and $\Delta^{17}\text{O}$ values of chondrules for individual chondrite groups. Data are from [9, 11, 13-14, 17-18].

SUMMARY

-Highly unequilibrated R3 clast NWA 753-IX contains a variety of chondrules with Mg#'s of 100-67.

-The petrologic subtype is estimated to be 3.15-3.2, suitable for detailed oxygen isotope study.

-More R3 clasts need to be examined.

R3 Data ??