

**IMPACT-GENERATED HYDROTHERMAL ALTERATION ON EARLY MARS IN PRESENCE OF CO<sub>2</sub>.**

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**Introduction:** Impact-generated hydrothermal systems may have caused widespread alteration of the crust on early Mars. We previously explored the range of alteration products using several Martian meteorites as crustal proxies: LEW88516 [1-3], Dhofar 378, Chassigny [4], and Nakhla [5]. Overall, the dominating alteration phases are oxides, hydroxides, clay minerals, and other hydrous silicates. Our results match well with the hydrous silicates (nontronite, chlorite) observed by OMEGA and CRISM [e.g., 6,7]. The two most recent findings by CRISM are carbonates [8] and serpentine [9]. While the latter is formed in our calculations from all olivine-rich starting compositions (LEW88516, Chassigny), carbonates require CO<sub>2</sub> (or HCO<sub>3</sub><sup>-</sup>, CO<sub>3</sub><sup>2-</sup>) in the system. To evaluate the effect of CO<sub>2</sub>, we conducted new calculations with ALHA84001 [10] and maskelynite [11] as starting compositions by exposing them to the same brine used before and, in a second set of calculations, with the addition of 0.5 mol H<sub>2</sub>CO<sub>3</sub>. We are using ALHA84001, because its secondary carbonates [10,11,12] provide a unique opportunity to compare model and mineralogical data. Moreover, it is the only sample of the ancient Martian crust [12] within the suite of Martian meteorites.

**Results and conclusions:** At a water to rock ratio (W/R) of 1 (P=110 bar, T=150 °C) the ALHA84001 alteration assemblage is dominated by serpentine and amphibole with chlorite, which is very similar to hydrous alteration of LEW88516. At intermediate W/R (~1000) the assemblage is hematite (51 %), nontronite, and quartz with pyrite; and at high W/R (~100,000) hematite dominates the precipitate. If 0.5 mole H<sub>2</sub>CO<sub>3</sub> are introduced into the brine, the assemblage at W/R=1 is still dominated by serpentine and amphibole, but ~5 % of carbonate are present. At higher W/R, carbonate dominates the precipitate (71 and 99 % at W/R = 1000 and 100000, respectively). If maskelynite reacts with a Fe-, Mg-, and Ca-bearing brine at W/R=1, the alteration assemblage is dominated by albite (37 %), mica, and zeolite with 4.6 % of calcite. At W/R=1000 the assemblage is kaolinite (56 %), siderite, and quartz. At very high W/R, it contains 95 % siderite. While assemblages at low W/R remain similar to carbonate-poor precipitates, high W/R can produce a carbonate-dominated assemblage. Furthermore, the calculations show that a carbonate-dominated assemblage can be formed from a Fe- and carbonate-carrying brine reacting with maskelynite. The calculations suggest carbonate abundances and compositions on Mars have the potential to provide estimates of W/R in subsurface hydrothermal systems.

**References:** [1] Schwenzer, S.P., Kring, D.A. 2008. *LPSC XXXIX*: #1817. [2] Schwenzer, S.P., Kring, D.A. 2008. *Workshop on the Early Solar System Bombardment*, #3027. [3] Schwenzer, S.P., Kring, D.A. 2008. *Workshop on Martian Phyllosilicates: Recorders of Aqueous Processes?* # 7014. [4] Schwenzer, S.P., Kring, D.A. 2009. *LPSC, XL*, #1421. [5] Schwenzer, S.P., Kring, D.A. 2009. *Workshop on Modeling Hydrous Environments on Mars*, #4015. [6] Bibring, J.-P. et al. 2005. *Science*, 307: 1576–1581. [7] Mustard, J.F. et al. 2008. *Nature*, 454: 305–309. [8] Ehlmann, B. et al. 2008. *Science*, 322: 1828–1832. [9] Ehlmann, B. et al. 2009. *LPSC, XL*: # 1787. [10] Gleason, J.D., et al. 1997. *GCA*, 61: 3503–3512. [11] Kring, D.A. et al. 1998. *GCA*, 62: 2155–2166. [12] Treiman, A.H. 1998. *Meteoritics*, 33: 753–764.